

**CORRELATION BETWEEN THE VOLUMES OF SLURRY AND
THE VOLUME OF FRACTURE IN BARNETT SHALE**

A Thesis

Presented to

the Faculty of the Department of Earth and Atmospheric Sciences

University of Houston

In Partial Fulfillment

of the Requirements for the Degree

Master of Science

By

Olawale John Adejuyigbe

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Olawale Adejuyigbe

APPROVED BY:

Dr. Evgeni M. Chesnokov, Professor
(Committee Chair)
Department of Earth and Atmospheric Sciences
University of Houston

Dr. Gennady M. Goloshubin, Professor
Department of Earth and Atmospheric Sciences
University of Houston

Renato Leite, Ph.D.
Senior Petrophysicist II
Petrobras America, Inc.

Dean
College of Natural Sciences and Mathematics

Dedication

This work is dedicated to my mother, the late Bosede Adejuyigbe who stood by me and offered invaluable moral and financial support all through my life.

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I sincerely acknowledge and thank my project supervisor, Dr. Evgeny Chesnokov, who allowed me to pursue this research, for his immeasurable input and guidance which helped in the completion of this work. The lessons learned from him will be an invaluable asset in my career. It would have been impossible to complete this work without his helpful supervision.

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Abstract

Microseismic fracture mapping has revealed that large fracture networks can be generated in many shale reservoirs and create multiple of plan complex structure. The ultra-low permeability in Barnett shale requires hydraulic fracturing to generate wellbore contact with the reservoir in order to achieve economic gas rates by the concept of a single fracture half-length. This is the reason for using fracture half-length of the hydraulic fracture area dimensions through the total fracture network length to estimate the volume of hydraulic fracture from a Dollie Thorell IV well in Barnett Shale. All data were released by Devon Energy for the purpose of this study.

As a way to better evaluate the effective producing network of the shale reservoir, Surfer VI and discrete bin method were used to estimate of the volume of hydraulic fracture. The results of the analyses from Surfer VI were from (i) Trapezoidal rule; and (ii) Simpson's rules, which gave an average result of 6.7×10^{10} cubic ft, the discrete bin method was calculated from the equation of fracture network half-length where length ($2x_f$) and width (x_n) represented the fracture spacing in the shale reservoir. The discrete bin method analysis gave a total result of 6.7×10^{10} cubic ft; this is the estimated figure for the volume of hydraulic fracture in very tight reservoirs. The correlation between the volume of slurry and the volume of the hydraulic fracture will establish the crack density. The graphic result of the correlation between the volume of slurry and the estimated volume of fracture showed that the crack density increases when the volume of the fracture increases. The crack density is always less than 1 due to leak-off of fracturing fluid into the reservoir during the hydraulic fracturing treatment.

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CHAPTER 1: INTRODUCTION

Studies of the volume of hydraulic fractures provide information about the untreated reservoir sections that can be targeted on subsequent treatments, as well as see whether a fracture intersects an aquifer or other geo-hazard. The volume of hydraulic fractures keeps the fracture propped open throughout the created fracture area and provides sufficient conductivity contrast to accelerate flow to the wellbore through the permeable pathway. This study focuses on the best method to calculate the volume of hydraulic fractures by comparing the volume of the slurry to the volume of the fracture. The micro-seismic location data and fracture treatment data from Dollie Thorell well in the Barnett Shale were integrated to calculate the volume of hydraulic fracture.

The Barnett Shale is among the most important onshore natural gas fields in North America and also the biggest in Texas. Billions of cubic feet of natural gas are produced from a great number of wells each year. The advancement in techniques for gas recovery and substantial gas value has caused recovery activity to increase dramatically within the last decades. The Barnett Shale is an ultra-low permeability reservoir that should be comfortably fracture-stimulated in an effort to achieve commercial production. Major fracture network operations are required for virtually any ultra-low permeability shale reservoirs with the purpose of improving the well performance. Microseismic fracture mapping is the main process that normally monitors the development of the fracture network.

1.1 Geologic Setting

The Barnett Shale of the Fort Worth Basin is classified as a homogenous undifferentiated black shale. Marine shale serves as the source for many different conventional clastic and carbonate reservoirs, which include the Pennsylvanian Bend and Strawn groups of the Fort Worth Basin (Pollastro et al. 2007). The Barnett Shale is now assumed to be a completely self-contained petroleum system, in which it serves as the source, seal and reservoir rock (Jarvie 2005). The Barnett Shale is known as one of the most prolific shale gas plays with unconventional resources. It produced more than 2.6 TCF (trillion cubic feet) of gas in 1987 (Durham 2007). Research on wireline logs as well as cores focuses on the sedimentology, internal geometry, and cyclical depositional patterns in the Fort Worth Basin (Figure 1.1) (Frantz and Jochen 2005). The Fort Worth basin is a north-south elongated Basin in north-central Texas, encompassing roughly 15,000 square miles (24,100 sq km). It is one of seven late Paleozoic basins associated with the Ouachita fold and overthrust belt (Meckel, et al. 1992). The basin is bordered on the north by the Red River Arch and on the south by the Llano Uplift and toward northeast by the Muenster Arch see (Figure 1.2). The complete history of the Fort Worth Basin reflects the Wilson cycle of the opening and subsequent closing of an ocean basin (Walper, 1982). The continental breakup of the North American part of Laurentia in the early Paleozoic gave

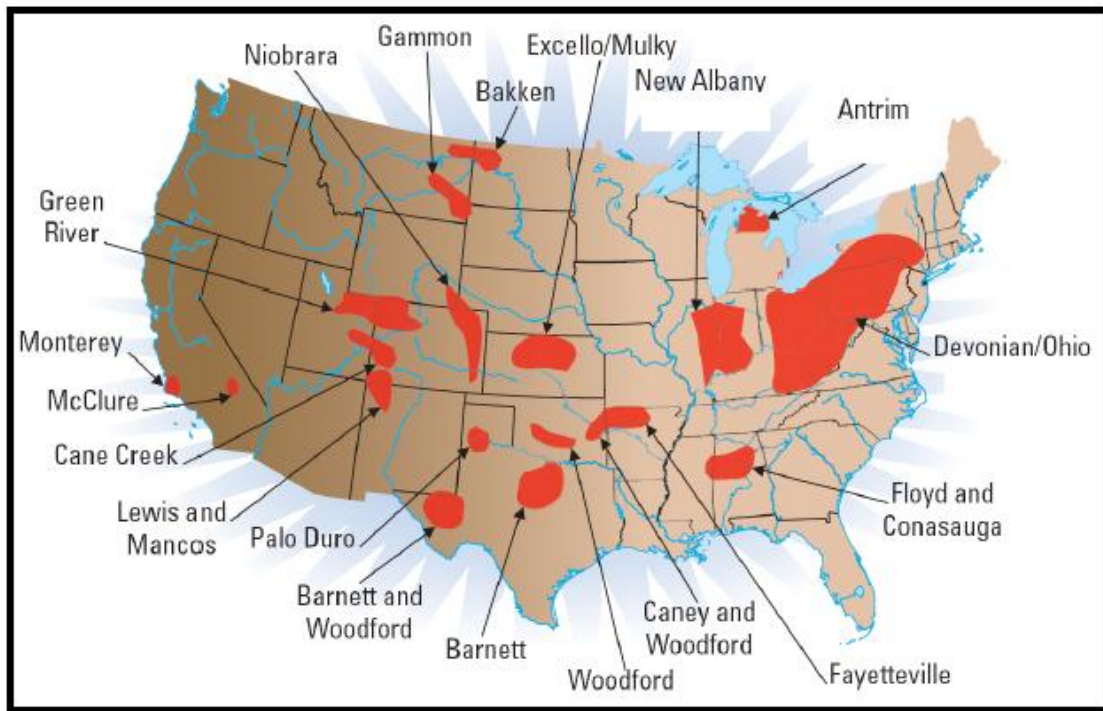


Figure 1.1 Major gas shale basins in the United States (After Frantz and Jochen, 2005)

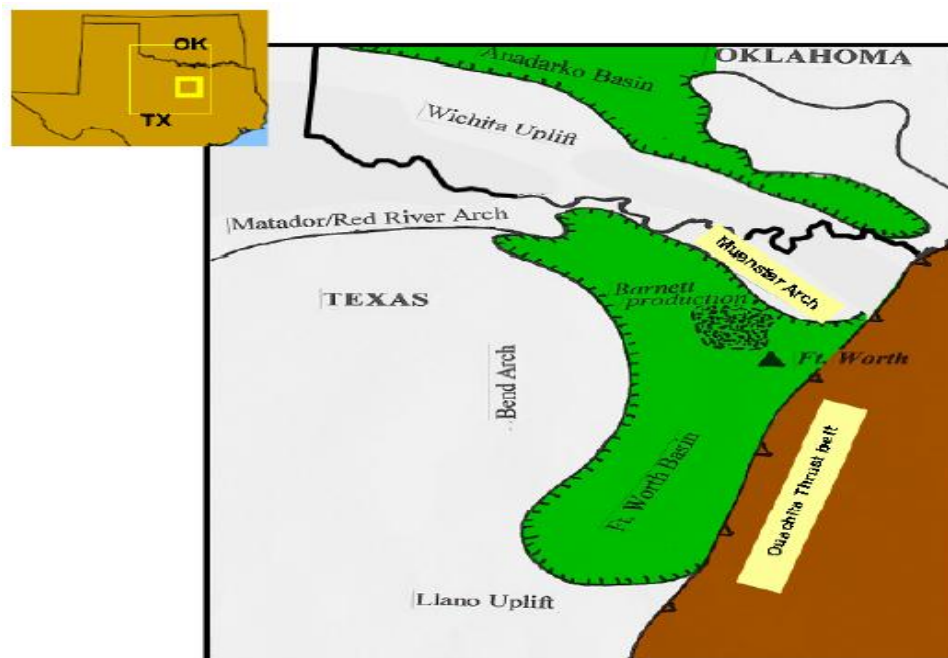


Figure 1.2 Regional setting of Fort Worth Basin modified (after Bawker, 2008)

rise to several aulacogens in nearby areas (Walper, 1977). As the margin of southern North America subsided and retreated from a spreading ridge, the Paleozoic sea, the Iapetus Ocean advanced. This invasion of the warm tropical sea is marked by widespread carbonate shelf deposition such as the Ellenburger Group of rocks. By Early Ordovician time the Iapetus Oceanic lithosphere began to subduct beneath the eastern margin of North America to form a marginal ocean basin between them (Walper, 1982).

In the Middle Paleozoic, there was a period of tectonic activity that resulted in erosion in the Fort Worth basin area marking the onset of the closure of the marginal ocean basin and reversal of the subduction polarity. With the onset of closure of the continents to form Pangaea, the formation of the Ouachita fold belt began, which represents the suture zone of the collision of the North America and the Afro-South American plates. The Ouachita foldthrust belt grew and evolved as the subduction complex - composed of the Ouachita facies strata which were deposited in the marginal ocean basin - was scraped from the subducting oceanic crust (Walper, 1982).

The subduction complex continued to be underthrust by the North American plate which made the outer arch migrate westward to make the hinge line to continue to retreat in the cratonic interior. A widespread marine transgression resulted in the formation of younger strandlines towards the craton suggesting that dormant aulacogens (zones of structural weakness) were reactivated by the subsequent collision of the North and South American plates, to eventually subside and become inundated by the sea (Hoffman, et al. 1974). By late Paleozoic time, the Fort Worth foreland basin formed in front of the advancing Ouachita structural belt. The basin was submerged and the Barnett

Shale and Comyn Limestone formations were deposited over the area of the Fort Worth basin (Walper, 1982). (Figure 1.3).

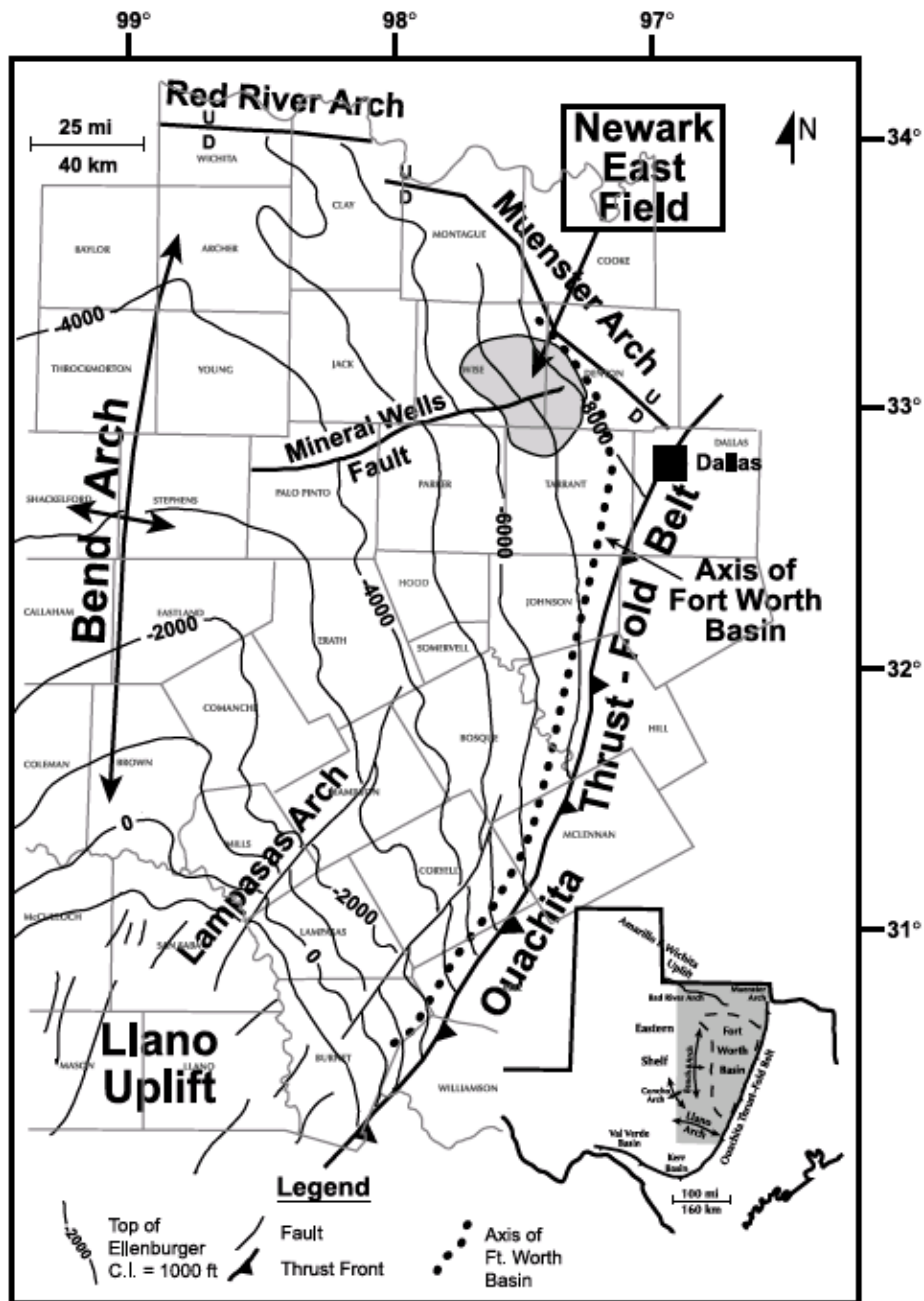


Figure 1.3 Regional setting and structure contour map of the Fort Worth basin, north-central Texas (Scott et al. 2005)

1.2 Scope of Work and Objectives

The aim of this thesis is to estimate the volume of hydraulic fractures by using micro-seismic events. The volume of hydraulic fractures is expressed in terms of the number, size and orientation of micro-seismic events. Knowing the volume of hydraulic fractures is good for correlation parameters for well performance. However, an important fracture is not the only parameter that drives well performance. There are other parameters such as fracture spacing for a given conductivity which can affect production acceleration and ultimate recovery (Fisher, et al. 2002). The largest fracture networks in the Barnett Shale show initial relationships between treatment size, network size, shape, and production response (Fisher, Maxwell and Urbancic, 2002).

Our results from micro-seismic fracture mapping show that network size is related to stimulation treatment volume. Figure 1.4 shows the large network which implies

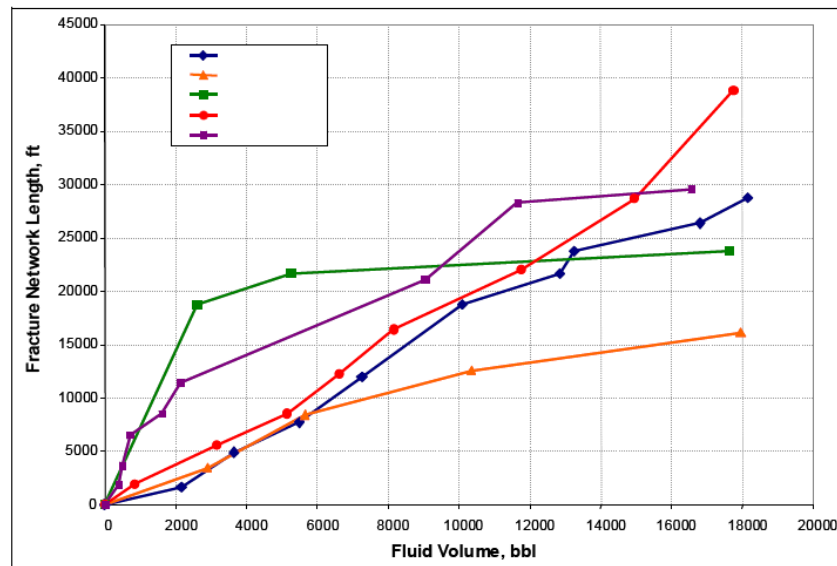


Figure 1.4 Relationship of total fracture network length as a function of total job fluid volume pumped (from Mayerhofer and Warpinski, 2008).

multiple fractures prompts the relationship between treatment volume and fracture network size. Millions of dollars are spent for fracture stimulation, but it has been difficult to evaluate and design stimulation treatment.

1.3 Previous Work

This study is part of an ongoing work aimed at better understanding the volume of hydraulic fractures as a correlation parameter for well performance. The size of the created fracture network, as indicated by the volume of hydraulic fractures, can be approximated as the 3-D volume of the micro-seismic event cloud. The estimation from micro-seismic mapping data is related to total injected fluid volume and well performance. A fluid mixed with chemicals called a “pad” is pumped to initiate the fracture and to establish propagation. This is followed by a slurry of fluid mixed with a propping agent which continues to extend the fracture and carry the proppant deeply into the fractures. After the materials are pumped, the fluid chemically “breaks back” to a lower viscosity and flows back out of the well, leaving a highly conductive propped fracture for oil and/or gas to flow easily from the extremities of the formation into the well (Gidley et al., 1989).

CHAPTER 2: HYDRAULIC FRACTURING AND MICROSEISMIC PRINCIPLES

2.1 History and Development of Hydraulic Fracturing

The first research to increase the production of commercial gas was performed in the Hugoton Gas Area in 1947 on Kelpner Well 1. As a fracturing fluid, a gasoline-based napalm-gel was used. However, these unpropped treatments did not increase production, leading to the belief that hydraulic fracturing did not represent any improvement in well performance (Gidley, et al., 1989). In 1949, 11 out of 23 wells where hydraulic fracturing treatments had been applied reported a significant increase in productivity in fields located in Wyoming, Colorado, Oklahoma and Texas (Clark, 1949). In 1952, the use of refined and crude oils as fracturing fluids gained popularity. Due to the lower cost, greater volumes per job were achieved. Nevertheless, a gradual change began in 1953 when aqueous-base fluids began being used, and by the end of 1963 up to 60 per cent of the fracturing jobs used this type of fluid (Hassebroek and Waters, 1964). The regular cost for a hydraulic fracturing treatment method ranged from \$2500 to \$3500 per well. Wells that had been hydraulically fractured proved to be more productive in comparison with those in which acid treatments had been used. (Patterson 1957)

In 1957, mathematical connections between overall fracturing performance, injection rate, and pumping time, as well as fracture size, were designed to predict fracture length. This provided the reason why some fluids were more effective and even why some pumping rates produced more effective end results. From this period on, fracturing treatments transformed from an experiential basis to technically- based

perspective (Hassenbroek and Waters, 1964). In the 1970's high tech massive hydraulic fracturing (MHF) treatments making use of great volumes of water started to be used. During these treatments as much as 1 million gallons of water and 3 million pounds of proppant were utilized. At that time, MHF was the main technique to economically improve tight reservoirs (Veatch, 1983). The earliest commercial vertical Barnett Shale well was the C .W. Slay No .1 drilled and stimulated in 1981 by Mitchell Energy. The initial fracturing productivity consisted of titanium, and zirconium-based crosslinkers with hydroxypropyl guar (HPG) and carboxymethyl hydroxypropyl guar (CMHPG). Treatments varied greatly: from 650-750 gallons of fracturing fluid to one million gallons and nearly about three million pounds of propping agent, at that time 25-30% of the whole U .S. oil reserves had been economically produced using this method (Veatch, 1983).

Hydraulic fracturing was preferred as a technique to optimize productivity in low-permeability reservoirs as well as to overcome damage in moderate- and high-permeability reservoirs. Hydraulic fracturing furnishes increased control of fines in unconsolidated formations; pressure decreases because of production is distributed over the surface area of the created surface, as opposed to the surface texture of the wellbore or gravel pack radius. This circulation of pressure generates a reduction in flow rate per unit area, which decreases flow velocity. A reducing of flow velocity will reduce formation fines movement (Parker, et al. 1994). In 1995, Union Pacific Resources (UPR) carried out the first "waterfrac" in the Cotton Valley Formation. A waterfrac is implemented by pumping large volumes of water with minimal amount of chemicals such

as surfactants and friction reducers. Normally 20/40 Ottawa sand is blended at 0.5 ppg. Early results showed a very similar production compared to that generated by former fracturing techniques but with 30-70% less fracturing costs. In 1997, using this technique represented about \$4.5 million in savings (Walker et al., 1998). After the effective usage of waterfracs in the Cotton Valley Formation, Mitchel Energy commenced experimenting in 1997 with waterfracs in the Barnett Shale. It was thought that similar accomplishments would be obtained with this particular treatment, and therefore different types of the treatment were implemented until a working design was reached. A current average fracturing job in the Barnett Shale only requires 750,000 gallons of slick water and 80,000 pounds of proppant pumped at 60 bpm with proppant concentrations from 0.1 to 0.5 ppg (Fisher, et al. 2002).

2.2 Microseismicity

Microseismic theory is analogous to that used in earthquakes; when fluid is injected into the formation causing changes in pore pressure, those changes affect the stability of planes of weakness, such as natural fractures and bedding planes. Shear slippages are produced in planes of weakness leading to failure and these are known as microseismic events. Microseismic activity in field applications has moment magnitudes which are measured based on the strength of the source from 3.5Ms to 2.5Ms. An earthquake usually needs to have a moment magnitude of starting time (t_3) to be felt at the surface (Warpinski, 2009). The receivers that record these events are an array of sensors located either on the surface or down-hole in an observation cell (Warpinski, 1998).

Analysis is performed to locate the source of each microseism and map the geometry and orientation of hydraulic fractures. There are two approaches to analyze the recorded data. The first approach uses the arrival time of P, and S-waves and formation velocities at different receivers located at different observation wells. The location of the microseismic event is then triangulated until the calculated location matches the observed arrival times. The second method involves the use of a single observation well with a multilevel array of receivers. Since more information is required, this technique requires higher-level technology receivers. The arrival times from the P- and S- waves, and the particle motion of the P-wave are required to estimate the azimuth of the microseismic event with respect to the position of the receiver (Warpinski, 1998).

2.2.1 Microseismicity Development

Microseismicity study has its origins in the 1930's when L. Obert and W.I. Duval discovered that a stressed rock pillar emitted microlevel sounds at a deep hard rock's mine (Obert, 1975). Since then several researchers have performed laboratory and field experiments to develop this technique; Obert and Duvall (1942, 1945-a, 1945-b) performed several experiments both in the lab and in the field, where they showed that different types of rock under compressive load generate acoustic emissions. Acoustic emission (AE) rate increased as the load was increased. The results of laboratory experiments on metals known as the Kaiser effect, noticed the effect of sample stress history on the production of acoustic emissions (Tam and Weng, 1995). Goodman (1963) also observed the relationship of stress state and AE rate during cyclic loading experiments that were performed on sandstone and quartz diorite samples. Other

researchers performed experiments under uniaxial and triaxial compressive stresses and found that the acoustic emission rate increased significantly as the compressive failure stress was reached (Barron, 1969, 1970; Mogi, 1962; Suzuki, et al. 1964; Mae and Nakao, 1968; Scholz, 1968a, 1968b).

During the 1970's, AE studies were underway in several geo-technical areas. Barron (1970) used a device to detect microseismic activity in specimens under triaxial load conditions. Haimson and Kim (1977) and Khair (1977) performed cycling uniaxial and triaxial compressive experiments in which AE were used to study fatigue mechanisms. Byerlee and Lockner (1977b) carried out hydraulic fracturing experiments in which AE was used to map the fractures; Lockner and Byerlee (1977a) also used AE to map fractures created during deformation of rocks under confining stress. During the 1980's and 1990's several researchers performed many laboratory experiments in which different materials were hydraulically fractured and AE was used to map the created fractures and study the fracture mechanisms (Majer and Doe, 1986; Matsunaga, et al. 1993; Van Dam, et al. 1998; Groenenboom, et al. 1999; Kranz, et al. 1990). Microseismic data have also been used more recently to estimate formation permeability, as reported by Shapiro, et al. (2006), Dinske, et al. (2010), and Grechka, et al. (2010).

2.2.2 Microseismicity Tools

Transducers used in microseismic applications are devices consisting of a coil around a mass hanging on a spring surrounding fixed magnets (Figure 2.1). Whenever the case (housing) moves, the coil moves with regards to the permanent magnet. The relative velocity is converted into an electromagnetic field. This magnetic field generates power

voltages that can be measured. Three-component receivers, where each axis is perpendicular to the opposite two, are used to record the polarization details from the P- and S-waves. Before actual recording can be attained, calibration of the receivers is performed with the intention of identifying the orientation and polarity response of the transducers; calibration is carried out by generating a source at a known location: this is known as a perforation shot. Generally, two of the axes are placed in sideways positions while the third is aligned in a vertical direction. Orientation of the receiver is achieved by using the location of the “artificial” source and the signals recorded by the two horizontal channels and the corresponding hodogram, which is a plot of the polarization of the signal (Sleepe, et al., 1995).

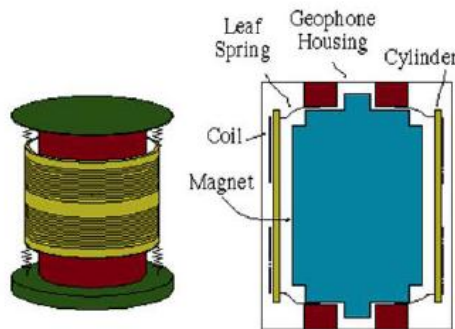


Figure 2.1 Schematics of Transducers (Barzilai, et al. 1998)

In microseismic applications, accelerometers are more suitable than geophones because of the many limitations of geophones. Geophones provide unsubstantial coupling to the borehole over an entire frequency band. Geophones are paired to the borehole via a locking arm that runs and clamps the unit to the borehole. The conventional swing-arm clamp typically includes a resonance around 200-400 Hz (Warpinski, et al., 1998); this

implies that at those frequencies the motions of the clamping unit usually do not go along with the motions of the borehole wall structure, causing a weakened coupling. Also, geophones present spurious modes at frequencies about 25 times higher than the natural frequency of the geophone. This is caused when the spring reaches a resonant frequency in the direction perpendicular to the axis of the geophone (Faber and Maxwell, 1996). Accelerometers, however, are built to have resonances above 2000 Hz, and they never contain spurious resonances.

Accelerometers are usually responsive, in comparison with geophones, at higher seismic frequencies because of their smaller electric noise (Sleeve, et al., 1995). Another essential criterion in selecting the best tool is the noise options; the signal-to-noise ratio (SNR) should be as large as possible in an effort to pick up weaker signals. When the intensity of the event cannot be enhanced, the noise should be decreased to attain a satisfactory SNR. According to Warpinski, et al. (1998), there are basically two forms of noise: cultural and electrical. Cultural noise includes that which forms as a result of gas bubbling via perforations from a region below the receivers, as well as truck and pumping activity. It is difficult to decrease this noise, although restorative action may be carried out to mitigate it. Electrical noise, on the other hand, is due to the transducers, electronic components, noise in the power supply, and pickup over unshielded components and wires (Warpinski, et al., 1998). The electrical-noise floor of the accelerometers is more advanced than geophones at frequencies challenged in microseismic services. This implies electrical noise of accelerometers is substantially

lower than that of the geophones at high frequencies (above 1000 Hz), allowing accelerometers to detect weak signals even at those frequencies (Sleeve, et al., 1995).

2.2.3 Microseismic Event Location

Microseismic events produce acoustic waves that travel throughout the earth. The location of a microseismic event is referred to as an inverse problem where the data are the arrival times recorded by different sensors, and the unknowns are the spatial coordinates (x, y, z) and the origin time of the microseismic event (Stein and Wyssession, 2003).

To describe how the inverse problem technique is used, let us imagine a MS event with unknown position $x = (x, y, z)$ and unknown origin time t_0 ; the arrival time varies between the travel time and the origin time:

$$d_i = t_0 + t_i \quad \text{Eq. 2.1}$$

t_i = the travel time, and this can be expressed as functions of spatial coordinate the sensor and the source:

$$t_i = \sqrt{\frac{(x_i - x)^2 + (y_i - y)^2 + (z_i - z)^2}{v}} \quad \text{Eq. 2.2}$$

where v is the velocity. The arrival time can be written in the simplest case of constant velocity as

$$d_i = t_0 + \sqrt{\frac{(x_i - x)^2 + (y_i - y)^2 + (z_i - z)^2}{v}} \quad \text{Eq.2.3}$$

This issue can be stated as data vector d which is from the function, A , acting on a vector m which represents the model factors, or unknowns

$$d = A(m) \quad \text{Eq. 2.4}$$

Simply, the inverse problem can be seen as given observed arrival times, and a model must be constructed to suit the observation. The procedure begins with a preliminary guess of the model (spatial coordinates and origin time). The initial guess will allow determining information, or appearance times, which we evaluate with the real noticed data, m . Usually the starting model generated wrong results; therefore, changes in the beginning style are necessary to arrive at a better remedy.

$$m_j^1 = m_j^o + \Delta m_j \quad \text{Eq.2.5}$$

Here Δm_j is the difference of the j th model parameter modification which makes a better fit of the recorded data. The information does not rely linearly on the model parameters; therefore, a linearization of the issue is necessary. This is acquired by expanding the data in a Taylor series about the starting model m^o and maintaining only the linear term.

$$\Delta d_i^o \approx d_i^o + \sum_j \frac{\partial d_i}{\partial m_j} \Delta m_j \quad \text{Eq. 2.6}$$

Therefore, the vector d_i^o contains arrival time calculations by using the initial guess of the model. Equation 2.6 is expressed in terms of the difference between the noticed data (arrival times) and those expected by the model.

$$\Delta d_i^o \approx \sum_j \frac{\partial d_i}{\partial m_j} \Delta m_j \quad \text{Eq.2.7}$$

The term $\frac{\partial d_1}{\partial m_j}$ is determined as the partial mixture matrix, G_{ij} ; therefore, the change

between the observed and calculated arrive time at a given model can be

$$\Delta d_i^o \approx \sum_j G_{ij} \Delta m_j. \quad \text{Eq. 2.8}$$

The model vector has four unknowns: source time to the three spatial coordinates; therefore, j varies from 1 to 4. However, i varies from 1 to n , where n is the variety of number of arrival time record by n sensors, which is usually greater than 4. Since the variety of lines and the variety of content in G are not equivalent, the matrix is not square, so it cannot be inverted. To resolve this issue equation, 2.8 has to be multiplied by GT , or the transposed of the partially derivative matrix, leading to

$$G^T \Delta d = G^T G \Delta m, \quad \text{Eq. 2.9}$$

and

$$\Delta m = (G^T G)^{-1} G^T \Delta d. \quad \text{Eq. 2.10}$$

This is the least square solution to modify in the model so the spatial coordinate and the origin time approach the observed data. We begin with an initial model m^0 , which is made up by an initial guess of the spatial coordinate and the source time. The recently expected arrival times are calculated and compared to the observed ones. The complete squared misfit is then measured as,

$$\sum (\Delta d_i^o)^2 = \sum (d_i - d_i^o)^2. \quad \text{Eq. 2.11}$$

The first model usually does not provide an excellent fit to the observed data; therefore, a further modification is necessary. The partially derivative matrix around the starting model can be approximated as

$$G_{ij} = \frac{\partial d_1}{\partial m_{j \ m^0}} \quad \text{Eq. 2.12}$$

Once the partial derivative matrix is calculated, the change in the model Δm^0 and a new model can be estimated to repeat the process:

$$m^1 = m^0 + \Delta m^0 \quad \text{Eq. 2.13}$$

The procedure recurs until the complete squared misfit is acceptably small.

Some signals have a low signal-to-noise rate due to the noise environment. This requires a solution so each sensor is given a certain significance according to its quality. This solution is known as the weighted least square and is expressed as

$$\Delta m = (G^T W_d G)^{-1} G^T W \Delta d \quad \text{Eq. 2.14}$$

Where W_d is the data weighting matrix and is obtained by inverting the variance-covariance matrix of the data, or vector.

2.3 Importance of Hydraulic Fracturing Technology

Study of the tracking of the fracture provides us with better knowledge of the fracture geometry, azimuth, connection, solidity, and length. This allows upgrades to be made to the fracture system in real-time, such as variations to push stress and rate or proppant capability. Moreover, the operator can recognize fluid activity patterns, fracture growth, connection, compaction, and whether the fracture and proppant are remaining "in zone" or "out of location." These crucial ideas allow the stimulation treatment plan to be enhanced, and provide ideas useful for long-term upgrades to the well space strategy, the well design, and the completion design. Ultimately, hydraulic fracture mapping applied

through microseismic tracking helps to boost the stimulated reservoir volume while enabling enhanced production rates and bookable supplies. Seismic waves go through attenuation (a decrease in the wave amplitude due to loss in energy), dropping, refraction, expression and function alterations by the time they are recorded by the receptors. The immediate compressional trend (P-wave) is the first stage to reach the indicator.

Fracture mechanisms and the frequency content of each signal help in understanding the fracture morphology. However, under similar conditions, tensile, compressive and more complex failure mechanisms have been observed in hydraulic fractures (Matsunaga, et al., 1993; Sasaki, 1998; Ishida, et al. 2004b). Microseisms do not map exactly where individual hydraulic fracture planes are located, but rather form an ellipsoid around the fracture, outlining the length, height, and azimuth of the fracture (Warpinski, 1994). Because microseisms may happen several feet off to one side of the fracture, no details on fracture size are accessible. The reason for this is most certainly the errors associated with hypocenter locations.

2.3.1 Hydraulic fracturing experiments at laboratory scale

Many researchers in recent years have performed hydraulic fracturing analysis at the laboratory scale. Haimson and Fairhurst (1969) tested hydrostone samples loaded polyaxially to simulate the three tectonic stresses encountered in the subsurface. They determined the type of fracture, its inclination and orientation and correlated those to the stress state in every sample. In addition, they determined the pressures at which the fracture triggered at different pressurization rates and different borehole sizes and

compared those values with those generated from theoretical criteria. Solberg, et al. (1977) conducted hydraulic fracture studies on triaxially stressed samples of oil shale and Westerly Granite. They obtained results from those samples with different stresses: higher than 29000 psi failed by shear, while samples with different stress lower than 29000 psi failed in tension.

Lockner and Byerlee (1977) performed many hydraulic fracturing experiments in Weber sandstone samples in which they used AE to locate the microseismic events. The samples were put through various stresses and fluid injection rates. Medlin and Masse (1979) chose four different varieties of limestone quarry rocks (Carthage, Indiana, Lueders, and Austin) for hydraulic fracturing laboratory experiments with the purpose of calculating fracture initiation pressure and orientation. Both cylindrical and spherical cavities were tried. Fracture initiation pressure was determined for each rock category making use of different injection fluids (non-penetrating grease and penetrating vacuum pump oil) under different conditions (hydrostatic stress and ambient) to monitor the effect of each parameter. End results of fracture initiation pressure were measured against those estimated by poroelasticity theory. Fracture initiation pressures produced from laboratory experiments are based on the poroelasticity theory over several ranges of hydrostatic stress, and vary depending on the rock properties. Medlin and Masse (1984) chose the Mesaverde sandstone, and the Carthage, and Lueders limestones for laboratory hydraulic fracturing analysis. They fractured the rock samples applying distinct injection fluids, injection rates, and confined stress state to estimate additional variables, including

fracture width, fracture length, and propagation pressure to calculate crack propagation theories.

Majer and Doe (1986) hydraulically fractured 300x300x450 mm salt blocks triaxially filled to compare the impact of confining pressure on breakdown pressure and the time dependency of breakdown pressure. Also, they applied AE to locate the microseismic events, as well as to investigate the behavior of the hydrofracture process. Cheung and Haimson (1989) conducted laboratory hydraulic fracturing experiments on fractured Niagara dolomite under a triaxial conditions. They investigated the conditions that regulate whether newer hydraulic fractures are produced or whether pre-existing fractures are reopened every time fluid is injected into the sample, and checked the results with values obtained theoretically. The results suggested that generation of new fractures or the reopening of pre-existing ones might be estimated in most circumstances, especially once the penetration of the injection fluid into the rock is considered.

Haimson and Zhao (1991) carried out many experiments hoping to explain the effect of borehole size and injection rate on the hydraulic fracturing break pressure, using granite and limestone samples. They concluded that borehole size and pumping rate influence breakdown pressure in the laboratory, but that impacts are insignificant in the field scale. Matsunaga, et al. (1993) performed laboratory hydraulic fracturing experiments in acrylic resin blocks, Inada granite, Komatsu andesite and Akiyoshi marble samples, using both water and oil as fracturing fluids. AE monitoring was used to interpret the fracturing mechanism during the fracturing process and the reaction of the fluid used. From focal mechanisms analysis, it was noted that events for all three rock

samples were the result of shear failure, while fractures in the acrylic resin blocks was the result of tensile failure. Masuda, et al. (1993) performed laboratory experiments on Inada granite samples in which they monitored the AE generated when fluid was injected into the samples. They planned two different experiments: one in which the sample was dried out and subjected to hydrostatic stress, and the second in which the sample was saturated and put through a differential stress. In the first experiment no AE was observed; however, in the second experiment, from the moment the water pressure was increased microfracturing was induced.

Van Dam, et al. (1998) performed hydraulic fracturing experiments on 30 mm (1 inch) cubic blocks of various substances. They made use of plaster, cement paste, and diatomite. Every single block was triaxially loaded to simulate in-situ stresses. AE was helpful to calculate the fracture radius as well as size of the non-penetrated area. They concluded that the variety of fracture radii after shut-in induces the leak-off volume. Song, et al. (2001) carried out hydraulic fracturing experiments on Table rock sandstone. They set out to identify whether this particular treatment (hydraulic fracturing) was easy for calculating in-situ stresses in highly permeable rocks. After assessing a couple of samples and several experimental parameters they saw a relationship between breakdown pressure and far-field stress. Song and Haimson (2001) researched the end results of pressurization rate and pore pressure on breakdown pressure using Table rock sandstone, and confirmed a relationship between breakdown pressure and the far-field stresses. Good agreement between experimental results and theoretical prediction of breakdown

pressure was found in the case of variable pressurization rates (Detournay and Cheng, 1992).

However, if pore pressure varied, the equivalent theoretical concept required some modification. Lhomme, et al. (2002) conducted hydraulic fracturing experiments at laboratory scale on Colton sandstone samples. A variety of fluid viscosities and injections rates were normally used to study fracture propagation fracture observations. In contrast to previous studies, they noticed that initiation pressure and breakdown pressure are not controlled by rates of pressurization or fluid viscosity. When they used high viscosity fluid and low injection rates they discovered a monotonic pressure drop after breakdown; whenever low viscosity fluid was used and injected at high rates, numerous fluid pressure rises were detected after the first pressure maximum. De Pater and Dong (2007) conducted several laboratory experiments to investigate the effects of confining stress and fluid rheology on hydraulic fracturing treatments using loose sand samples. The injection fluids used in the experiments were highly viscous Newtonian fluids (500 Poise), a bentonite slurry, a cross-linked gel, and the related cross-linked gel with fine quartz particles. Confining stresses, starting from 29 to 29000 psi, were used. Only when cross-linked gel with quartz particles was applied as the injection fluid was a hydraulic fracture created at all confining stresses used, and at any pumping rate.

Surdi, et al. (2010) conducted hydraulic fracturing experiments on a pair of identical quartz-rich, Carbon Tan sandstones, and observed the acoustic activity. To simplify fracture initiation two diametrically opposed slits $\frac{1}{4}$ -inch in length were cut. They also modeled the distribution of stress focusing in the sample within well pressurization; they

associated stress concentration during loading and fracturing with the localization of acoustic emissions in space and time. It is most evident that many researchers have tried to understand the hydraulic fracturing process through controlled experimental results.

2.3.2 Current Status of Hydraulic Fracturing Technology

Hydraulic fracturing was first established as a strategy to prevent well damage and increase production in conventional and tight gas reservoirs. However, growth in interest in unconventional reservoirs has taken hydraulic fracturing to new technological frontiers. The prospective advantages of refracturing have found the interest of oil and gas operators for more than five decades. If an original treatment is inadequate or the existing proppant deteriorates over the time, re-fracturing the well re-establishes linear flow into the wellbore (Dozier et al., 2003). Wells with effective initial treatment can be re-stimulated by creating a new fracture that propagates along a different azimuth than the original fracture which exposes more net pay to the wellbore (Dozier, et al., 2003). The key to successful development of an unconventional gas reservoir is to create complex fracture networks that contact a large reservoir volume (Mayerhofer, et al., 2006). However, the nature and degree of the fracture complexity must be understood in order to select the best stimulation strategy (Cipolla, et al., 2010).

Enhanced hydraulic fracturing techniques attempt to make the stimulation process more effective. One of these is hydraulic fracturing using carbon dioxide (CO₂). Verdon et al. (2010) studied a hydraulic fracturing field case in which water and supercritical CO₂ were injected under similar conditions. Microseismicity was used to observe performance variations. The microseismic event location showed similar styles in both

cases, but in the case of water injection the fracture extended further laterally. When CO₂ was injected, microseismic activity appeared far above the injection factor, indicating that its higher buoyancy permitted CO₂ to move from top to bottom. The magnitude of the events caused in both situations is identical. Verdon et al., (2010) determined that despite the change in compressibility, solidity and viscosity, both liquids (water and CO₂) have identical styles of caused seismicity.

2.4 Considering Anisotropy

Isotropic, linear, elastic materials are completely characterized by two independent constants, usually Young's modulus and Poisson's ratio. Therefore, calculation of both P-wave and S-wave velocities becomes independent of the direction of wave propagation (Mavko et al., 2003). In anisotropic rocks, elastic characterization is not as simple as the isotropic case; up to 21 elastic constants are needed to describe the elastic behavior of anisotropic rocks. We often encounter two types of anisotropy: transverse isotropy and azimuthal anisotropy. The former has a hexagonal symmetry with five independent elastic constants, where the symmetry axis is normally perpendicular to the bedding. The latter is caused by stress anisotropy. Azimuthally anisotropic rocks may have 5, 9 or 13 independent elastic constants, depending on stress orientation and the intrinsic properties of the rock (Wang, 2002).

Most crustal rocks are found experimentally to be transversely isotropic as a result of the preferred orientations of anisotropic mineral grains, the shapes of isotropic mineral and those of cracks or thin bedding of isotropic or anisotropic layers (Thomsen, 1986).

Thomsen further pointed out that in most cases of interest to geophysicists, the anisotropy is weak (<10 percent). He indicated that a weak transverse isotropic formation can be characterized by three anisotropic parameters (ϵ , δ , and γ) and two velocities. The equations of the phase velocity variation as function of the angle are as follows:

$$V_p(\theta) = \alpha_0(1 + \delta \sin^2 \theta \cos^2 \theta + \epsilon \sin^4 \theta), \quad \text{Eq. 2.15}$$

$$V_{sv}(\theta) = \beta_0(1 + \frac{\alpha_0^2}{\beta_0^2}(\epsilon - \delta) \sin^2 \theta \cos^2 \theta), \quad \text{Eq. 2.16}$$

$$V_{sh}(\theta) = \beta_0(1 + \gamma \sin^2 \theta). \quad \text{Eq. 2.17}$$

The angle between the wave front normal and the symmetry axis is known as the phase angle θ . α_0 And β_0 are the P-wave and S-wave velocities measured to the symmetry axis.

The magnitude of anisotropy in shale is greater than that estimated by Thomsen (Sondergeld and Rai, 2011); Sondergeld and Rai (1992) measured anisotropies as high as 42% in shear velocity and Hornby, et al. (1999) estimated P-wave anisotropy of 38%. Berryman (2008) proposed the following equations for strong anisotropy:

$$V_p(\theta) = \alpha_0 \left(1 + \epsilon \sin^2 \theta - (\epsilon - \delta) \frac{2 \sin^2 \theta_m \sin^2 \theta \cos^2 \theta}{1 - \cos 2\theta_m \cos 2\theta} \right) \text{ and} \quad \text{Eq. 2.18}$$

$$V_{sv}(\theta) = \beta_0 \left(1 + \frac{V_p^2(0)}{V_s^2(0)} - (\epsilon - \delta) \frac{2 \sin^2 \theta_m \sin^2 \theta \cos^2 \theta}{1 - \cos 2\theta_m \cos 2\theta} \right), \quad \text{Eq. 2.19}$$

where θ_m is the incidence angle near which the extreme SV-wave behavior occurs (Berryman, 2008) and is given as:

$$\tan^2 \theta_m = \frac{c_{22} - c_{44}}{c_{11} - c_{44}}. \quad \text{Eq. 2.20}$$

Any time a borehole is drilled in a transversely isotropic formation, the stresses around the wall of the borehole are not nearly as simple, almost suggesting the formation was isotropic. Aadnoy (1987) used equations to calculate the system stress around the borehole wall whenever a well is drilled in a transversely isotropic formation (Figure 2.2). The system stress calculation is split up into the components that determine the total hoop stress as show by

$$\sigma_{\theta\theta} = \sigma_{\theta\theta 1} + \sigma_{\theta\theta 2} + \sigma_{\theta\theta 3}, \quad \text{Eq. 2.21}$$

where $\sigma_{\theta\theta}$ the total hoop stress at the borehole wall , $\sigma_{\theta\theta 1}$ is the contribution due the borehole pressure, $\sigma_{\theta\theta 2}$ is the contribution of normal stresses, and $\sigma_{\theta\theta 3}$ is the shear stress contribution.

$$\sigma_{\theta\theta 1} = P_w \frac{E_\theta}{E_x} \left[k - n(\sin^2 \theta + k \cos^2 \theta) + (1 - k^2) \sin^2 \theta \cos^2 \theta \right], \quad \text{Eq. 2.22}$$

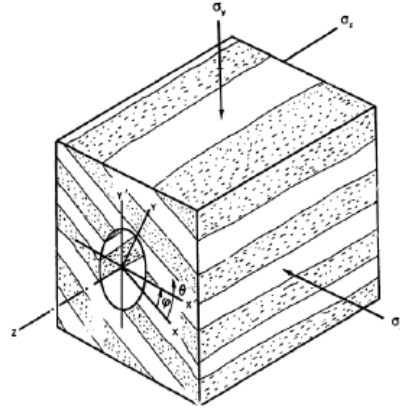


Figure 2.2 Geometry showing a borehole in a transversely isotropic medium where the X-axis is aligned with the direction of the bedding and the Y-axis is perpendicular to the bedding (Aadnoy, 1987). σ_x , σ_y , σ_z are the three principal stresses. σ_x and σ_y are aligned with the X'- and Y'-axis, respectively. Θ is the angle from the X-axis and ϕ is the angle measured from the bedding orientation to the horizontal axis (Castano, 2010)

where P_w is the hydrostatic pressure in the borehole, θ is the angle from the X-axis (from the axis of isotropy), E_θ is Young's modulus in the direction tangent to a position at a given angle, θ , from the direction of the bedding, E_x is the Young's modulus along the X-axis (parallel to bedding). K , n are given as

$$k = \left(\frac{E_x}{E_y} \right)^{\frac{1}{2}} \quad \text{and} \quad \text{Eq. 2.23}$$

$$n^2 = 2 + 2k. \quad \text{Eq. 2.24}$$

The value E_θ/E_x can be calculated from the following expression:

$$\frac{E_x}{E_\theta} = \sin^4 \theta + 2 \sin^2 \theta \cos^2 \theta + k^2 \cos^4 \theta \quad \text{Eq. 2.25}$$

The contribution of normal stresses is given as

$$\begin{aligned} \sigma_{\theta\theta} = & \sigma_x \frac{E_\theta}{E_x} \left[-\cos^2 \varphi + (k+n) \sin^2 \varphi \right] k \cos^2 \theta + \left[(1+n) \cos^2 \varphi - k \sin^2 \varphi \right] \\ & \sin^2 \theta - n(1+k+n) \sin \varphi \cos \varphi \sin \theta \cos \theta + \sigma_y \frac{E_\theta}{E_x} \left[-\sin^2 \varphi + (k+n) \cos^2 \varphi \right] k \cos^2 \theta + \\ & \sigma_y \frac{E_\theta}{E_x} \left[-\sin^2 \varphi + (k+n) \cos^2 \varphi \right] k \cos^2 \theta + \left[(1+n) \sin^2 \varphi - k \cos^2 \varphi \right] \sin^2 \theta - n(1+k+n) \sin \varphi \cos \varphi \sin \theta \cos \theta \end{aligned} \quad \text{Eq. 2.26}$$

where φ_x and φ_y are the principal stresses in the horizontal and vertical axes respectively. The angle φ is the angle measured from the bedding orientation to the horizontal axis as was shown (Figure. 2.1). The shear stress contribution is given by the following expression:

$$\sigma_{\theta\theta} = \tau_{xy} \frac{E_\theta}{2E_x} (1+k+n) \left(\left[-n \cos 2\varphi + (1+k) \cos 2\theta + k - 1 \right] \frac{\sin 2\varphi}{\sin 2\theta} \right), \quad \text{Eq. 2.27}$$

where

$$\tau_{xy} = 0.5(\sigma_y - \sigma_x)\sin 2\theta. \quad \text{Eq. 2.28}$$

2.4.1 Formation permeability from microseismic data

State of the art technology has permitted using microseismic information not only for applying to hydraulic fracture but also for the evaluation of in-situ formation permeability. For this purposes two different techniques are utilized: the r method and the inversion approach (Grechka, et al., 2010).

2.4.2 R-technique for permeability estimation

The r- technique depends on the spatio-temporal dynamics characteristic of caused microseismic clouds. The volume of the injected fluid must be equal to a sum of the fluid volume stored in the fracture and the fluid volume which goes into the formation. A straight planar height-fixed fracture is considered. Under these conditions the half-length, L , of the hydraulic fracture is given as a function of the injection time as the following (Economides and Nolte, 2003):

$$L(t) = \frac{Q_i t}{4h_f C_L \sqrt{2t + 2h_f w}}, \quad \text{Eq.2.29}$$

where Q_i is the average injection rate of the treatment fluid, h_f is the average fracture height and w is the average fracture width, t is the injection time, and C_L is the fluid-loss coefficient. In the case of the hydraulic fracturing of a low permeability formation, such as tight gas sandstones, the fracture body represents the main permeable channel in the

formation (Shapiro et al., 2006). The induced fracture changes the stress state in its vicinity, leading to the occurrence of microseismic events at a distance very close to the hydraulic fracture (Warpinski, 2000). Therefore, equation 2.30 can be considered as a one-dimensional approximation for the triggering front of microseismicity in the case of a penetrating hydraulic fracture (Shapiro, et al., 2006):

$$L = \sqrt{4\pi Dt} , \quad \text{Eq. 2.30}$$

where D is the apparent diffusivity and t is the injection time.

During hydraulic fracturing treatments the fracture growth is managed by the liquid loss effects. Generally, this means that the cumulative volume of the lost fluid is considerably larger than the volume of the hydraulic fracture. The fluid loss effects are controlled by the fluid-loss coefficient, C_L , which is characterized by the apparent diffusivity, D , (Shapiro et al., 2006):

$$C_L = \frac{Q_i}{8h_f \sqrt{2\pi D}} , \quad \text{Eq. 2.31}$$

Also, the fluid loss coefficient can be interpreted neglecting the near-surface effects (e.g., filter cake) from the pressure difference between the fracture and the far field reservoir ΔP . Economides and Nolte (2003) approximated the fluid-loss coefficient based on these assumptions to be

$$C_L = \sqrt{\frac{k_r \phi c_r}{\pi \eta_r} \Delta P} , \quad \text{Eq. 2.32}$$

where k_r is the in-situ reservoir permeability, ϕ is the formation porosity, C_r and η_r are the compressibility and viscosity of the reservoir fluid and ΔP is the difference between the

average injection pressure and the far-field reservoir pressure (Shapiro, et al. 2006).

Substituting Equation 2.31 into Equation 2.32 and solving for the formation permeability

$$k_{r-t} = \frac{\eta_r}{128\phi c_r D} \left(\frac{Q_i}{h_f \Delta P} \right)^2. \quad \text{Eq. 2.33}$$

Equation 2.33 is the in-situ formation permeability calculated using the r technique. The distance from the perforation point to the event location as a function of time is shown by the r plot. Equation 2.30 symbolizes the parabolic envelope that better explains the higher limit of a lot of microseismic activity, which is reliant on the apparent diffusivity value.

CHAPTER 3: MATERIALS AND METHODS OF STUDY

CHAPTER 3.1 General Statements

The unique low flow behavior of low permeability reservoirs is due to poor utilization of reservoir pressure and to reservoir heterogeneity. Poor utilization of reservoir pressure is caused in part by depositional issues like very small grains of detrital muds (clays); other factors include diagenesis related to clay precipitation, massive cementation, pressure compaction, etc. Reservoir heterogeneity is related to depositional and post-depositional (diagenetic) events. These include the vertical/lateral heterogeneity that shows the medium and large scale geological features for turbidite deposition and fault and differential diagenesis, which include the migration and hydrocarbon generation. These characteristics show that low permeability reservoirs are simply poor conductors of fluids.

Ultra-low permeability shale reservoirs require a major fracture network to maximize well performance. Microseismic fracture mapping has revealed that huge fracture networks are typically induced in many shale reservoirs. In shale reservoirs, anywhere complex network structures in different planes are formed, the ideas about that particular fracture half-length in addition to conductivity are insufficient to portray well performance.

Thus, reservoir permeability and reservoir heterogeneity are permanent constants that we sometimes cannot change. Although this is widely known, it is possible to modify our mechanism for accessing the reservoir (i.e., the well). We may also change our development method to ensure best performance and treatment of any reservoir. The

calculation of shale permeability remains problematic, particularly when considering the understanding of the end results of standard flow measurements (e.g., steady-state permeability measurements). Neuzil (1994) emphasizes less the special properties of shale permeability and more on the "regions" shown on a chart of porosity versus their logarithms of permeability. This context is important in recognizing that shales/clays have very high porosity and low permeability, which is important when considering trends. Revel and Cathles (1999) implement a power-law model for estimating permeability in shaly-sands making use of porosity and shale volume.

The influence of an in-situ stress field on a natural fracture or fault should be monitored in situ by its direction determined by that stress field. Evaluating the relationship of presented stress data to existing regional structures provides the ability to differentiate between natural fractures and new fractures being created. Although it is inconclusive to say that faulting or other geologic structures that are regularly visible on the surface near the area of a treatment well are an indication of the subsurface structure, at times surface features may work as useful signals of the valuable structural features in the subsurface. The response to an uncertainty of a new fracture, whether a new or pre-existing, was initiated has various impacts for the responsible failure mechanisms. With the utilization surface arrays and broad areal coverage laid arrays, declining mechanisms along fracture planes are often identified, giving the angle of the fault or fracture plane and also the path of slip, and details on the possible tensile area of the failure mechanism.

Microseismic monitoring surface arrays produce imaging coverage over the total stimulated volume of a treatment well. Simply because the span of areal coverage for the

several channels accepting event signals deals with a massive area, it is possible to investigate the first P-wave particle motion changes across the array and determine a fault plane connected with that motion (Figure 3.1) For nearly vertical faults, the first motions across the array will have zero P-wave amplitude radiated along a line parallel to a fault, and array stations will locate positive or negative motion on specific array areas.

This technique of interpretation is performed with surface based vertical component geophones and probably will not need the use of three-component geophones. The integration of this solution with information regarding the structural style in the area provides identification of the displacement on the fault as normal or reverse dip-slip without the need to do a full source mechanism inversion.

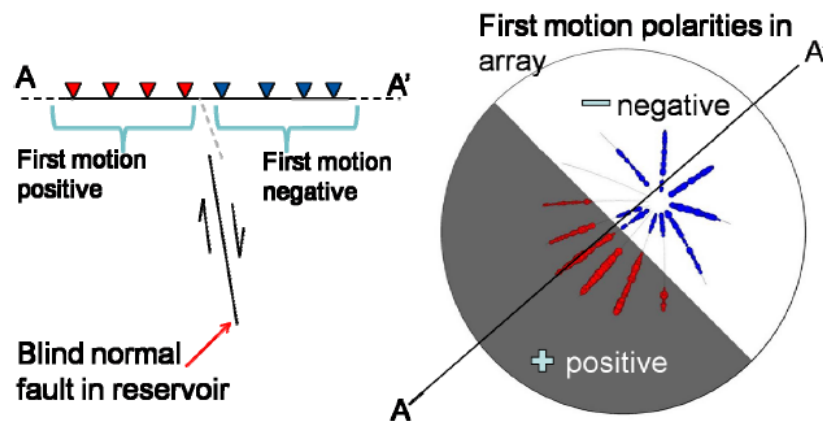


Figure 3.1. Example of surface-based microseismic geophone array configuration. The asymmetric layout of the receiver pattern optimizes signal for a horizontal well deviated to the southwest (Williams, 2009).

If perhaps a variety of strong events describe two intersecting styles about 30 degrees apart, it is clear that the treatment fluid located and reactivated two nearby fault planes (Figure 3.2.) The direction in which the fault planes slipped is not distinct from the

geometry, but by combining the geologic details with the wellbore information, the final nature of the failure mechanisms and fault interactions was identified. For events describing a 120° azimuth trend, a fault plane was described that parallels that 120° azimuth, which signifies that the first motions of the seismic energy P-waves to the geophones in the array were positive in one area on this nodal line and negative on the other side. This kind of result can come from dip-slip motion on a steeply dipping fault in the subsurface.

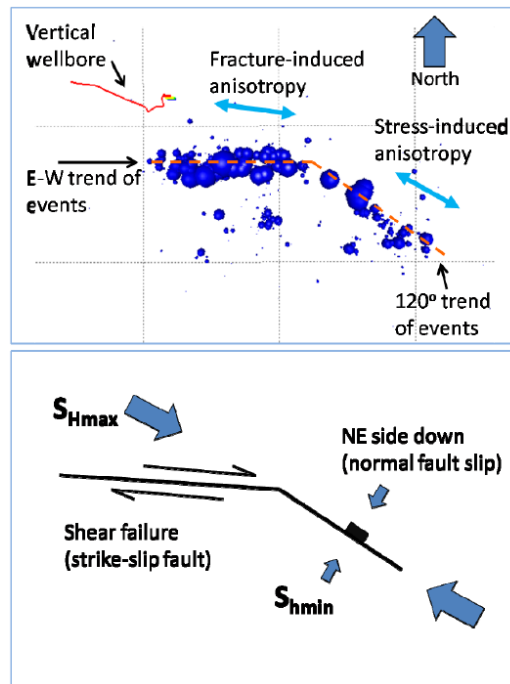


Figure 3.2 Failure mechanisms of two interacting faults reactivated by hydraulic fracture stimulation (Williams, 2009).

The events that describe the east-west pattern did not have this particular first motion of P-waves. The concept of the microseismic events is reactivation and possibly continuance of an east-west oriented fault in strike-slip motion; while it grew, it intersected a

southeasterly striking fault in order for normal dip-slip motion. The analysis is in accordance with the anisotropy measurements in the wellbore that determined maximum horizontal stress parallel to the strike of the normal dip-slip fault. Figure 3.3 shows a situation of a complex fracture network activated by the microseismicity of a stimulation treatment is useful to describe the upper bounds of fracture length, and their height and dip. An assessment of the event patterns with the subsurface structure means that the microseismicity has reactivated recent structures. The largest events align parallel to the strike of major faults unidentified in the subsurface in the interpreted seismic horizons. The most significant events create a direction parallel to the faults with orientation nearby to the maximum horizontal stress direction, and a secondary trend of weaker events seems to be forming at an angle to the first trend. The two orientations exist in the subsurface, the results are interpreted to reflect reactivated existing faults and fractures. The great microseismic events describe a trend at an angle over 25 degrees from the fault trend. In this case, the effective stress in the reservoir is east-northeast (roughly sub-parallel) to the faults, so faults in this specific orientation are supposed to be critically stressed, indicating it is in an orientation that is likely to slip (Zoback, 2007). Source mechanism inversions performed for this data set demonstrate principally dip-slip normal fault displacement on the fault planes oriented approximately 80° azimuth (Figure 3.3).

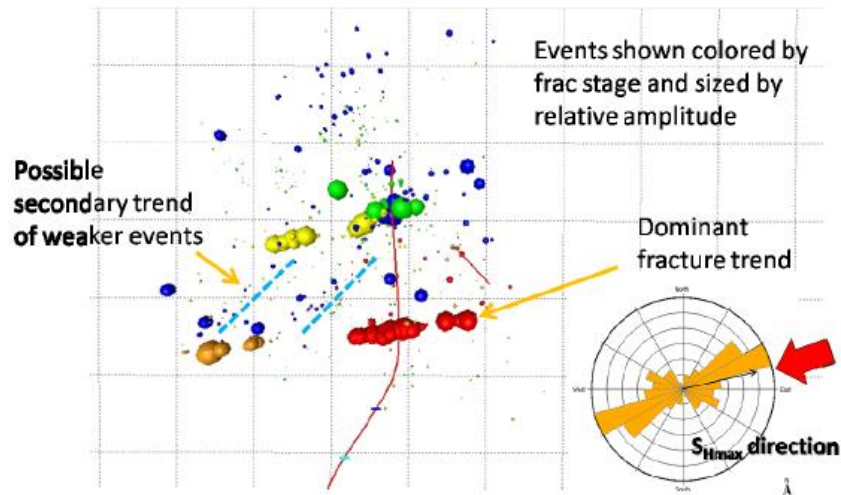


Figure 3.3 Microseismic monitoring result with some stages showing better defined trends than others. The trends are parallel to major faults mapped nearby in the subsurface. The failure mechanisms for the events are consistent with normal dip-slip motion along such faults (Williams 2009).

3.2 Microseismic Data

Data used in this study were obtained from Devon Energy Incorporated. The fracture treatment data consisted of the treating pressure, total slurry used, slurry rate, and proppant concentration. The micro-seismic location data consisted of the X, Y, and Z locations of all the events east to west and from north to south, time of location in seconds, epicentral distance, and number of usable P- and, S- wave arrivals. The slurry data are the number of microseismic events and the number of microseismic cubes (See Appendixes Table 1 and Table 2).

3.3 Methodology

The main objective of this study is to calculate the volume of hydraulic fracture from micro-seismic events and compare the result with the field measured volume of slurry. Two methods are suggested. The first method is the use of a software called Surfer 6 to plot and estimate the volume of hydraulic fracture. The second method is based on the use of the excel discrete bins technique for estimating the volume of hydraulic fracture.

3.3.1 Surfer application method

The workflow of the Surfer application started by creating grid on the X, Y, and Z locations data (Figure 3.4). The first task was to use the software to create an XYZ data file, which contained three columns of data values. The first two columns were the coordinates for X and Y data points and the third column was the Z. The XYZ data were entered into column ABC in an Excel spreadsheet (Figure 3.5) Surfer 6 is powerful software that has the ability to plot, contour or make surfaces. The grid files can produce a grid-based map and a grid report. The software created a grid from the data in columns XYZ where X, Y and Z represented Easting, Northing, and Elevation, respectively (Figure 3.6). Surfaces were made from the block diagrams generated by drawing lines representing the grid X and Y lines (the grid columns and rows). At each intersection of a column and row (at each grid node), the height of the surface was proportional to the grid Z value at that point. The number of columns and rows in the grid file determines the number of X and Y lines drawn on the surface in (Figure 3.7). The volume of the hydraulic fracture can be calculated from the generated surface.

The Surfer 6 application can also plot the XYZ information from the field data into 2D and 3D in a cube without changing the location data XYZ into surfaces. Although it was assumed that the best estimation could be calculated from the 3D plot because it showed the true projection of the micro-seismic event with respect to time and space as seen in (Figure 3.8a). The procedure of plotting the events in 3D used the XY, YZ, and XZ data only, so there were three different individual plots that were needed to estimate the volume of hydraulic fracture.

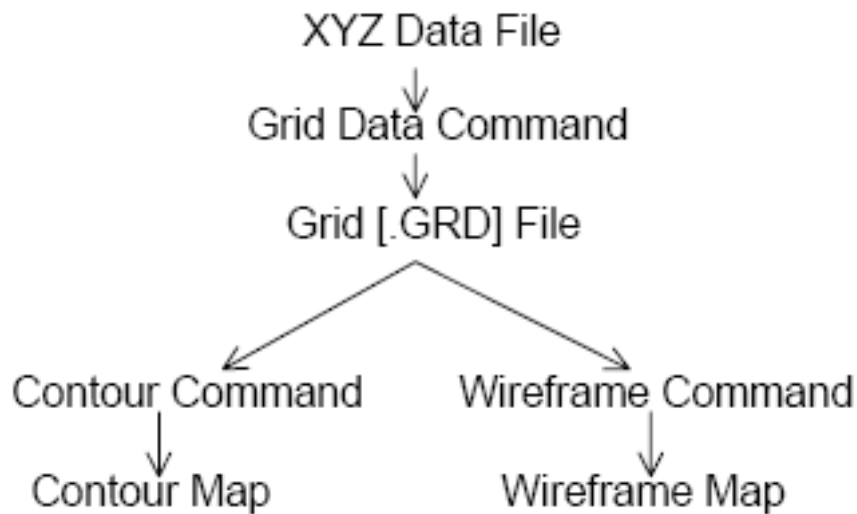


Figure 3.4. Flow chart illustrating the relationship between XYZ data, grid files, contour maps, and wireframes maps.

	A	B	C
1	X Data	Y Data	Z Data
2	0.1	0	90
3	9	3	48
4	1.3	7	52
5	4.7	1	66
6	1.7	5.6	75
7	6	1	50
8	2.5	3.6	60

Figure 3.5 The XYZ data placed in Columns A, B, and C in Surfer application.

Grid Data - C:\Program Files\Golden Software\Surfer8\Samples\TU... [?] [X]

Data Columns (47 data points)

X: Filter Data...

Y: View Data

Z: Statistics

☒ Grid Report

Gridding Method: Advanced Options...

Output Grid File:

Grid Line Geometry

	Minimum	Maximum	Spacing	# of Lines
X Direction:	0	9	0.09090909091	100
Y Direction:	0	7	0.09090909091	78

Figure 3.6 the grid data dialog box showing the gridding parameter to use when creating a grid file in Surfer application.

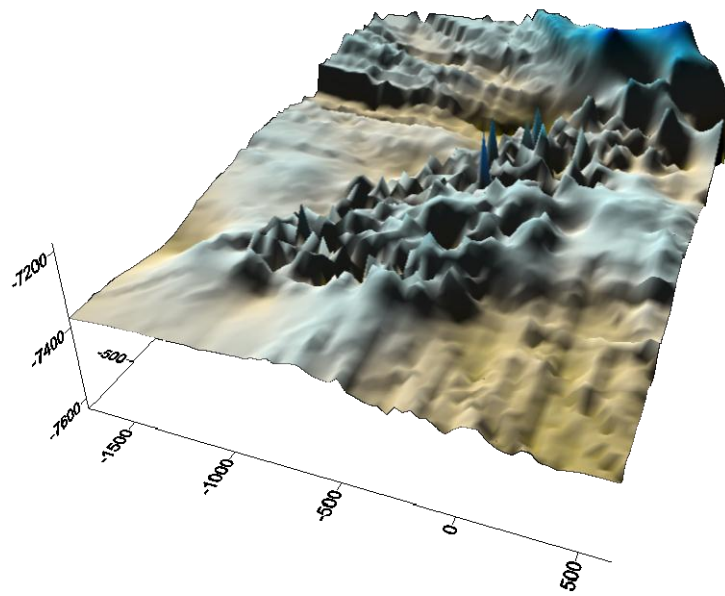


Figure 3.7 The surface from the Surfer X, Y, and Z parameters.

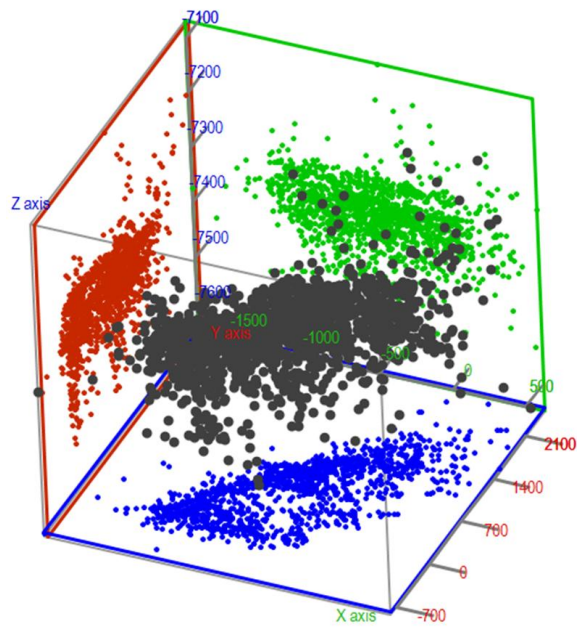


Figure3.8a Three dimensions of the micro-seismic events in Surfer.

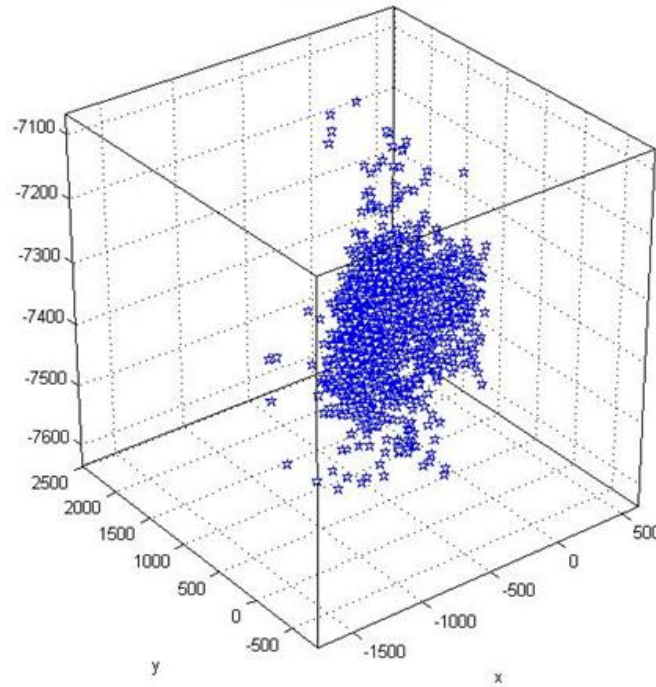


Figure 3.8b Three-dimensional projection of the all the microseismic events.

The entire Surfer plots were gridded into 100 by 100 in both axes. The volume of hydraulic fracture was estimated from the micro-seismic events by using 100 by 100 grids to form a rectangular box around the cloud. These rectangular boxes were measured from the maximum and minimum values in the x and y axes direction. The micro-seismic events were also plotted in 3D view in Surfer with the projection of the XY, XZ, and YZ at middle of the cube. Each point represented each coordinate from XYZ respectively, as show in (Figures 3.8a and Figure 3.8b). The estimated volume could be calculated from the event in the three -dimensional box of XYZ.

3.3.2 Excel Discrete Bins Method

This method is used to estimate volume of hydraulic fracture microseismic discrete bins. Constant width bins (e.g., 100ft wide) are drawn in the principal fracture direction from the borehole to the furthest event in the specific bin on both sides of the borehole either from the XY, XZ, and YZ. The individual bin areas are then summed up to approximate the total fracture area. The calculation volume of hydraulic fracture also requires an estimate of the fracture area network height in each discrete bin within the reservoir section. Figures 3.9a, 3.9b and 3.9c show the microseismic discrete bin method used for the Dollie Thorell well.

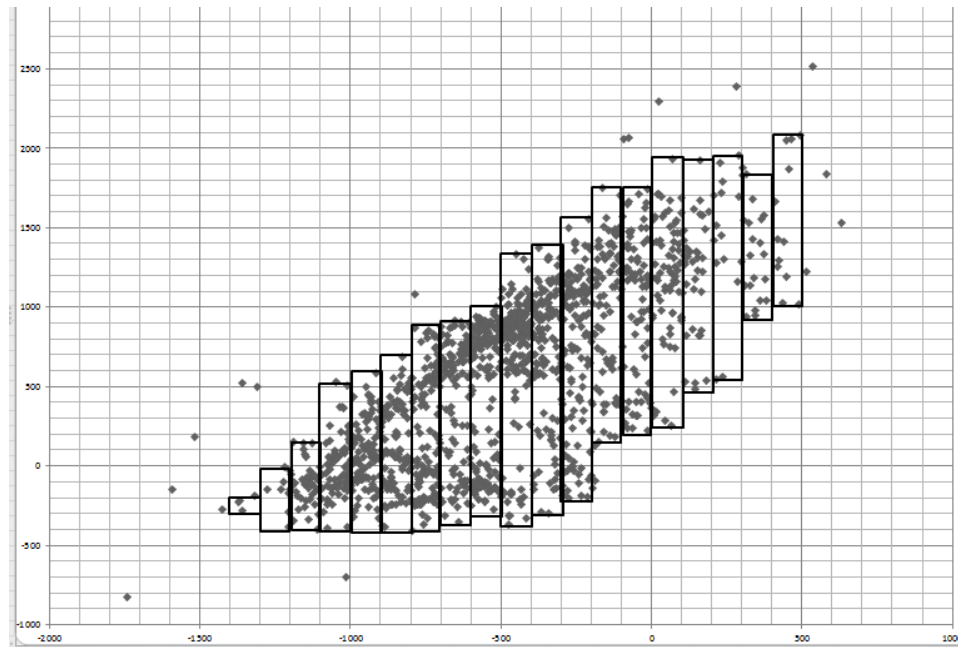


Figure 3.9a The XY axes of the micro-seismic event from the Discrete bin method.

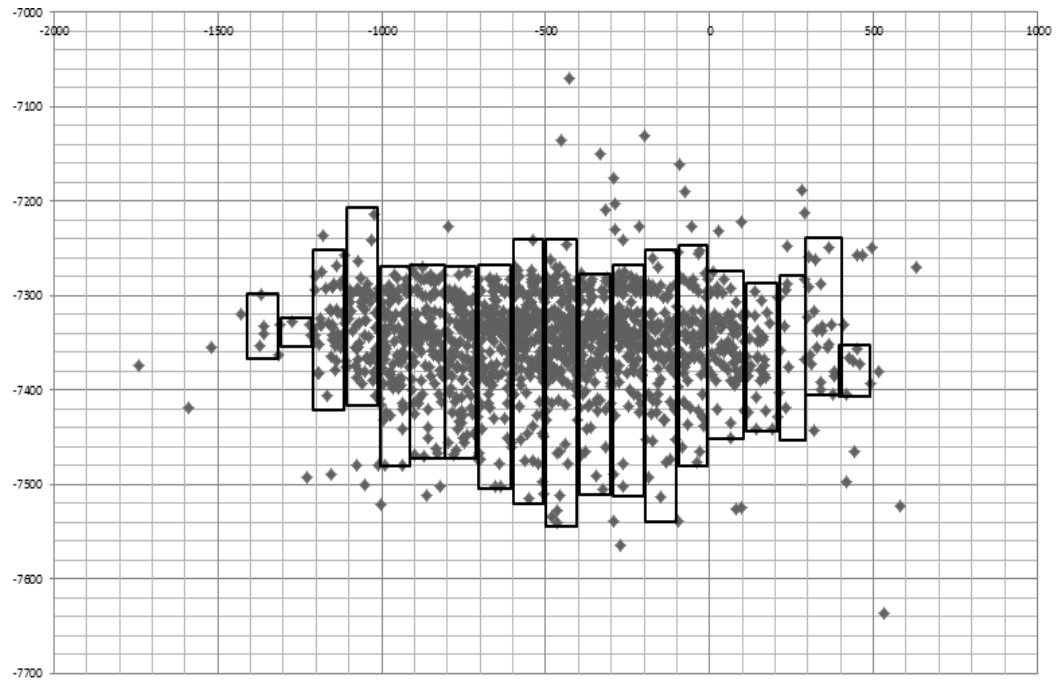


Figure 3.9b The XZ axes of the micro-seismic event from the Discrete bin method.

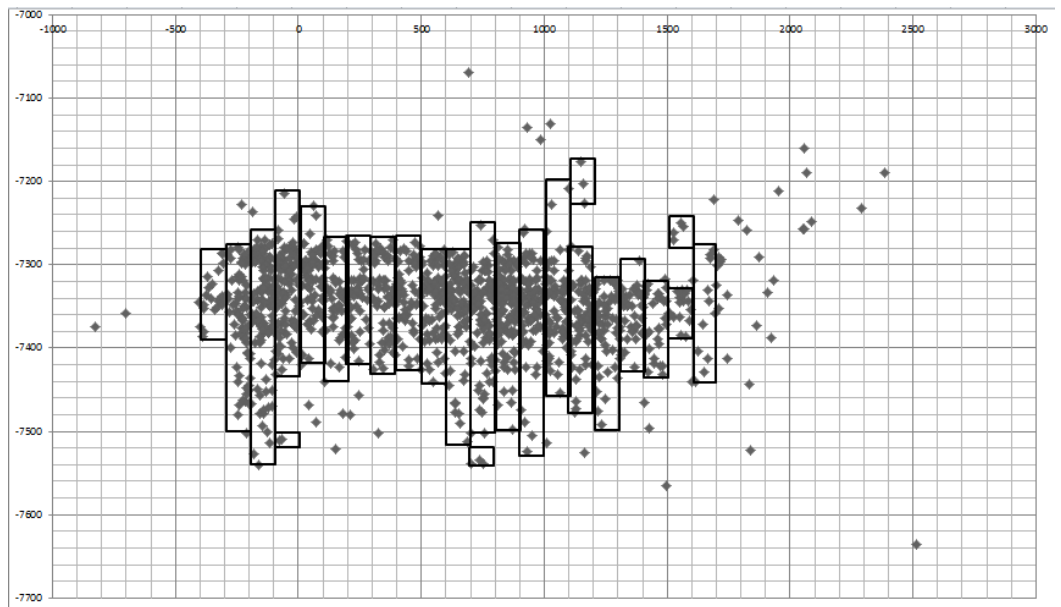


Figure 3.9c The YZ axes of the micro-seismic event Discrete bin method.

CHAPTER 4: RESULTS AND DISCUSSION

4.1 Surfer Application Result

Surfer 6 software was used for surface gridding and volume calculation. The input XYZ data (micro-seismic event data from the location) was available for the top of the Barnett Shale. No base of the interested interval was used. The way the code was written in this Surfer 6 program gridded the input XYZ data, then converted this gridded point of the XY into surfaces. The program calculated the entire area of the surface and multiplied it with the z values from the input data which is considered to be height of this event. The volume calculation thus involved the extent of the surface area and the height of the Z value of the input data. The Surfer 6 gave the output of the volume results using three rules: (i) Trapezoidal rule, (ii) Simpson's rule, and (iii) Simpson's 3/8 rules. The total volumes calculated are 6.7×10^{10} cubic ft, 6.7×10^{10} cubic ft and 6.7×10^{10} cubic ft, respectively. The results are also presented in Figure 4.1.

4.2 Excel Discrete Bins Method Result

This is the method by which the XY data (see Columns 1 and 2 in Table 2 in the appendix) were plotted and then gridded into different bin. A total of nineteen bins were constructed with intervals of 100ft in X- axis and 100ft in Y-axis, this procedure resulted in the formation of 19 different bins .

Grid Volume Computation

Upper Surface

Grid File Name *C:\Users\Seun\Contacts\Desktop\MSLocation.true.12.grd*

X Minimum: -999.59

X Maximum: 1742.7

X Spacing: 33.8×10^{12}

Y Minimum: -828.25

X Maximum: 2512.68

X Spacing: 33.7×10^{12}

Z Minimum: $-7.631.5 \times 10^{10}$

Z Maximum: 1742.7

Lower Surface

Level Surface defined by Z=0

Volumes

Z Scale Factor: 1

Total Volumes by:

Trapezoidal Rule: -6.7×10^{10}

Simpson's Rule: -6.7×10^{10}

Simpson's 3/8 Rule: -6.7×10^{10}

Cut & Fill Volumes 0

Positive Volume [Cut]: 0

Negative Volume [Fill]: 6.7×10^{10}

Net Volume [Cut Fill]: 6.7×10^{10}

Figure 4.1 Calculated volumes by using Surfer software Version 6 for the Barnett Shale. The surface height was used to define a lower limit for the volume definition.

Volume was calculated from the product of area (fracture spacing) and height. The height in Figure 4.2 is the difference between the highest and the lowest values from the discrete bin in relative to the XY axes (Chesnokov, 2009).

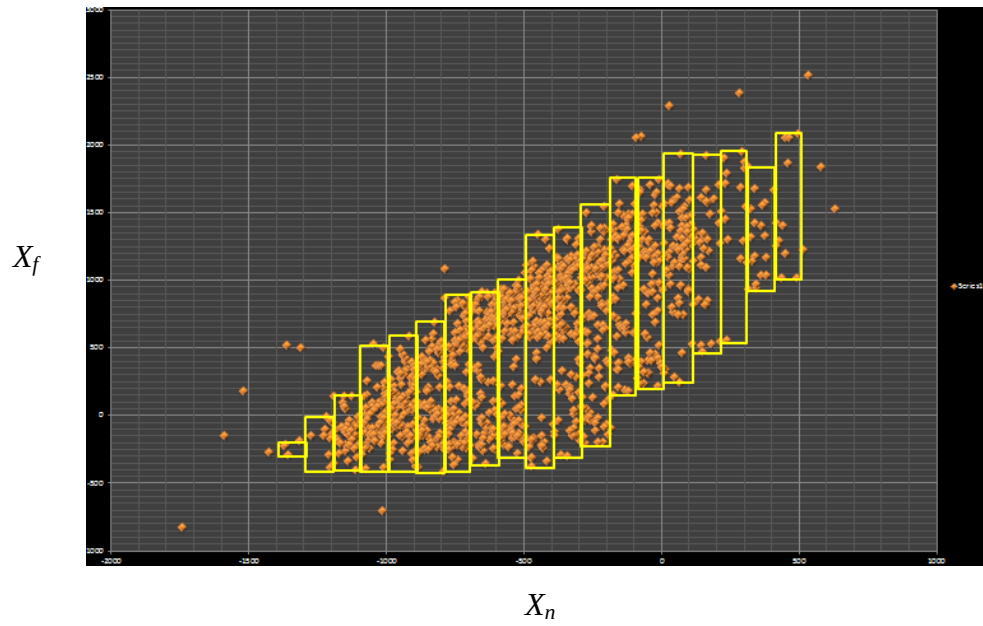


Figure 4.2 The fracture spacing of the microseismic events into different bins using the Discrete Bin method.

The volume of hydraulic fracture (V_f) calculated from this method was $6.7 \cdot 10^{10}$ cubic ft. All calculations are shown in a spreadsheet in Table 1 in the appendix. The calculated volume of fracture (V_f) was then correlated with the volume of slurry from the field (V_i) to determine the relation fracturing crack density η .

4.3 Discussion of Results

The purpose of estimating the volume of hydraulic fracture in this study is to reflect the nature and extent of any fractures that may have been created by the over-pressurization of the well. This is related to pore pressure changes which cannot move very far from the actual fracture planes, unless natural fractures in alternate directions are opened and hydraulically enhanced as a network structure thus serving as a conduit for fluid

movement. This means that a large micro-seismic event (cloud) structure must be approximately equivalent to the actual fracture network size. Thus, the microseismic event cloud structure detected by hydraulic fracturing provides a means of estimating the volume of hydraulic fractures in very tight reservoirs (Mayerhofer and Warpinski, 2008).

The total volume of the hydraulic fracture that was estimated from both methods provides details of the effectively producing fracture structure or spacing. Apart from reservoir modeling, which also provides an avenue to better evaluate the effectively producing network, the seismic signal characteristics also provide the source of the mechanical deformation that resulted in the micro-seismic event. Maxwell, et al. (2002) introduced a concept that could eventually be used to characterize fracture density. This concept was similar to the discrete bin method, which is the area of the hydraulic fracture shown in Figure 4.2. In contrast to conventional single fracture modeling using fracture half-length of the hydraulic fracture area, dimensions are given by the total fracture network length ($2x_f$) and width (x_n) (Figure 4.2). However, the key property for evaluating the volume of hydraulic fracture, is the total sum of all fracture network segments within the reservoir, which is a strong function of fracture spacing. The equation for the calculation of the total fracture length to estimate the volume of hydraulic fracture as a function of fracture network half-length x_f , width x_n and fracture spacing Δx_s is:

$$L_{total} = \frac{4x_f x_n}{\Delta x_s} + 2x_f + x_n . \quad \text{Eq. 4.1}$$

The relationship between the volume of the created hydraulic fracture, (V_f), the volume of the slurry (V_i) and the fracturing crack density (N_c) is given in Table 1 in the Appendix as:

$$N_c = V_f / V_i \quad \text{Eq.4.2}$$

Crack density is always less than 1 due to leak-off of fracturing fluid to the reservoir pore space during the hydraulic fracturing treatment. Figure 4.1 shows that the crack density increases when the volume of the fracture increase. The crack density is a function of orientation produced by the deviatoric stress field, which led to velocity differences and anisotropy. The fracture's anisotropic elastic constants can be described in terms of a fracture-compliance matrix, which can be added to the compliance matrix of the host rock to find the equivalent-medium elastic constants (Schoenberg, 1980; Schoenberg and Sayers, 1995). Crack density and aspect ratio can be related to the fracture compliance matrix for aligned fracture sets (Hudson and Liu, 1999; Liu et al., 2000), and the compliances of multiple fracture sets can be combined (Schoenberg and Sayers, 1995).

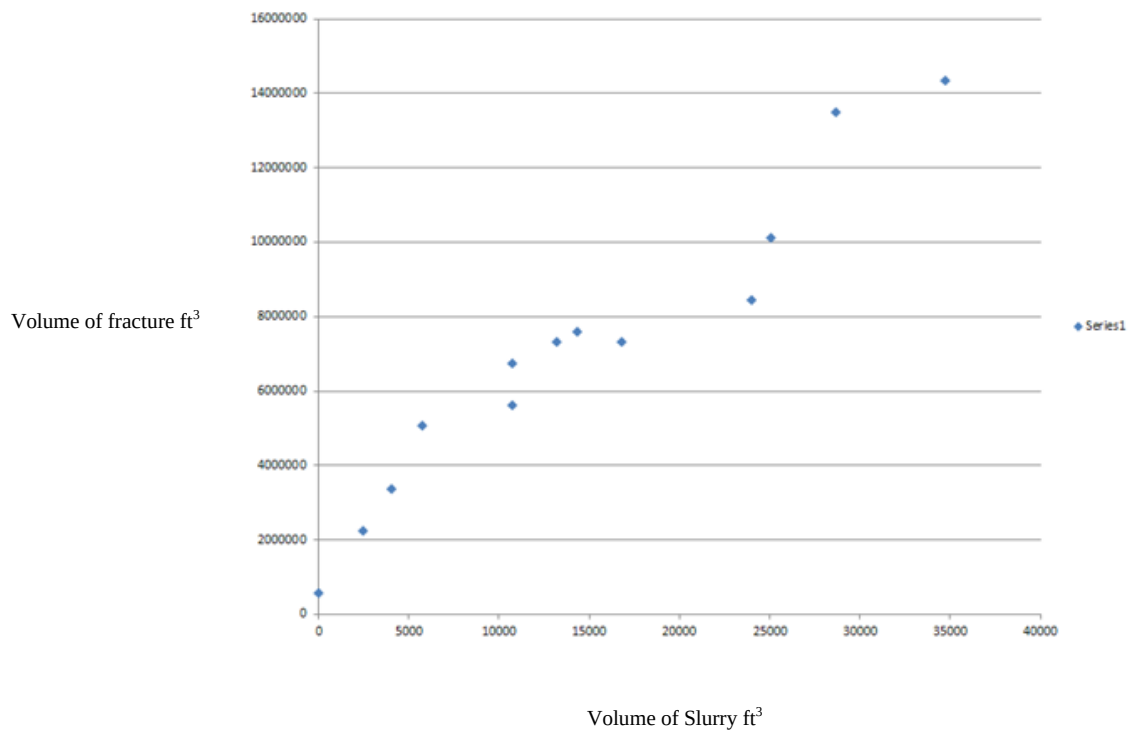


Figure 4.3 Correlation between volume of slurry and volume of fracture from table

APPENDIX 1

Area	xmin	xmax	ymin	ymax	Number of Events	Number of Cube	Cumulative Number of the of Cube	Area	Volume	Vol of fracture(m3)	Volume Slurry (m3)	Nc=Crack Density	Time(Sec)
1	-1400	-1300	-300	-200	3	2	2	10000	3750000	562500	0	5.33333E-06	299
2	-1300	-1200	-400	0	25	8	10	40000	15000000	2250000	2475	1.66667E-06	395
3	-1200	-1100	-400	200	41	12	22	60000	22500000	3375000	4051.5	1.82222E-06	977
4	-1100	-1000	-400	550	70	18	40	90000	33750000	5062500	5737.5	2.07407E-06	1321
5	-1000	-900	-400	600	81	20	60	100000	37500000	5625000	10720.5	0.00000216	1321
6	-900	-800	-400	700	95	24	84	120000	45000000	6750000	10720.5	2.11111E-06	1473
7	-800	-700	-400	900	102	26	110	130000	48750000	7312500	13213.5	2.09231E-06	1540
8	-700	-600	-300	1000	105	27	137	135000	50625000	7593750	14311.5	2.07407E-06	1692
9	-600	-500	-300	1150	103	26	163	130000	48750000	7312500	16804.5	2.11282E-06	1758
10	-500	-400	-400	1320	93	24	187	120000	45000000	6750000	17887.5	2.06667E-06	1910
11	-400	-300	-300	1400	39	11	198	55000	20625000	3093750	20380.5	1.89091E-06	1977
12	-300	-200	-200	1590	76	22	220	110000	41250000	6187500	21478.5	1.84242E-06	2129
13	-200	-100	100	1800	98	30	250	150000	56250000	8437500	23973	1.74222E-06	2195
14	-100	0	200	1800	101	36	286	180000	67500000	10125000	25054.5	1.4963E-06	2347
15	0	100	210	1960	82	34	320	170000	63750000	9562500	27547.5	1.28627E-06	2414
16	100	200	430	1940	104	48	368	240000	90000000	13500000	28647	1.15556E-06	2566
17	200	300	510	1980	58	28	396	140000	52500000	7875000	31140	1.10476E-06	2632
18	300	400	900	1820	80	40	436	200000	75000000	11250000	32223	1.06667E-06	2784
19	400	500	1000	2100	92	51	487	255000	95625000	14343750	34717.5	9.62092E-07	2851

Table 1. Result of the values used in calculating the volume of the fracture and crack densit

Table 2. The Micro-seismic Location Data (Input Data)

X, Y, Z locations relative to Monitor Well and other parameters associated with a microseismic event, including

Rxy - Epicentral distance = $\sqrt{X^2 + Y^2}$, ft

Nsp - Number of usable P-, S-wave arrivals x 1000,
Referred to the case of "Number of geophones" = 12.
Thus the limit superior = $1000 \cdot (2 \cdot 12) = 24000$

Tm= Tc-To = Difference between start time of the current
Event file and that of the 1st event file (in seconds)

O.time - Difference (in milliseconds) between the current events
Origin time and the event file start time

E.C. - The energy class as Brigg's logarithm of relative
Seismic energy ($E.C. = \log_{10} (E / E_{ref})$)

X, ft (W-E)	Y, ft (S-N)	Z, ft	Rxy, ft	Nsp	Tm, sec	O.time mls	E.C.
56.58	1322.37	-7348.85	1323.58	21081	2131	20.4	-3.4
-311.15	1262.99	-7345.44	1300.75	21753	2135	29.7	-3.3
-205.63	1362.94	-7379.39	1378.36	23408	2157	72.6	-2.3
-171.46	1063.61	-7454.34	1077.34	23162	2209	95.2	-3.1
-455.99	687.09	-7511.7	824.63	18121	2423	71.7	-3.8
74.13	1409.54	-7336.15	1411.49	21069	2467	13.4	-3.4
-929.54	407.26	-7364.68	1014.84	18785	2585	55	-4
-402.53	893.37	-7296.88	979.87	19354	2667	61.4	-4.1
-374.4	859.11	-7329.06	937.15	23782	2757	70.7	-3.5
-485.42	835.37	-7338.75	966.17	20349	2847	64.3	-3.8
-664.14	640.36	-7314.98	922.57	22821	2939	71.3	-3.3
-110.36	1318.88	-7329.43	1323.49	20282	2947	29.4	-3.2
-1018.92	109.6	-7378.58	1024.8	21027	3011	57.5	-3.4
-526.12	802.14	-7323.63	959.29	23532	3187	105.2	-1.9
-595.9	729.12	-7381.51	941.65	22764	3187	101.3	-2.4
-631.94	767.89	-7358.13	994.49	23578	3187	101.9	-2.3
-875.89	382.63	-7404.89	955.82	20151	3227	65.7	-3.5
-21.61	1138.05	-7332.1	1138.26	21834	3246	48	-3.1
-207.08	1383.29	-7336.44	1398.7	23009	3258	75.6	-2.6
65.52	1133.97	-7389.46	1135.86	18301	3288	44.4	-3.7
-929.9	59.38	-7343.35	931.79	20475	3360	63.4	-3.6
-738.82	559.79	-7353.08	926.94	18833	3386	61.6	-3.9
-636.48	700.4	-7501.65	946.4	22033	3486	59.8	-3.2
-253.12	944.89	-7365.06	978.21	21778	3486	67.1	-3.3
288.76	1158.36	-7367.55	1193.81	20498	3524	44.1	-3
-948.46	275.39	-7364.92	987.63	18169	3550	59.9	-3.7
-252.62	1357.18	-7349.31	1380.49	18214	3559	21.6	-2.9
-463.27	1108.28	-7277.89	1201.21	23322	3580	85.9	-1.9
-464.18	639.5	-7467.06	790.2	20287	3606	76.9	-3.5
-1067.48	252.96	-7380	1097.04	20664	3610	51.5	-2.7
-964.86	-79.53	-7397.78	968.13	20741	3626	54.3	-3.5
-927.37	353.85	-7288.42	992.58	18933	3641	61.8	-3.7
-388.95	810.91	-7369	899.36	20706	3647	269.5	-3.6
-518.43	743.55	-7365	906.44	19809	3647	70.1	-3.6
-183.69	1134.81	-7353.79	1149.58	19302	3647	40.9	-3.4
-214.26	1114.76	-7423.13	1135.16	21302	3655	48.2	-2.9
3.89	946.34	-7331.2	946.35	21978	3702	70.8	-3.1
-464.44	665.09	-7318.95	811.2	19530	3708	81	-2.9

-686.4	651.56	-7379.83	946.4	23790	3720	103.9	-2.9
-645.68	563.76	-7328.5	857.16	21262	3724	110.5	-2.5
-694.26	550.58	-7441.2	886.08	23742	3728	107	-2.5
-729.72	473.75	-7332.04	870.02	23025	3728	73.9	-3.2
113.1	1266.09	-7318.98	1271.13	21013	3760	33.6	-3.1
73.58	1167.26	-7340.07	1169.58	19519	3810	44.6	-3.1
-504.02	814.86	-7360.42	958.14	22614	3818	66.8	-3
-421.97	780.92	-7334.18	887.63	23263	3840	106	-3.4
-628.14	583.36	-7337.61	857.24	21899	3850	76	-3.5
-479.27	802.97	-7414.44	935.13	23506	3853	98.5	-2.1
-476.95	760.54	-7387.52	897.72	19718	3890	71.5	-3.8
-754.11	449.95	-7421.29	878.14	22052	3893	105.1	-1.4
-662.74	689.01	-7319	956.01	23327	3893	65.1	-2.6
-101.36	1110.43	-7342.06	1115.05	20490	3917	89.4	-2.6
-899	243.23	-7334.35	931.32	19923	3928	68.2	-3.4
-703.8	496.02	-7334.85	861.03	23282	3933	110.8	-0.6
82.94	1161.85	-7526.03	1164.81	22009	3965	41.4	-2.8
82.56	1542.42	-7332.29	1544.63	21773	3990	87.1	-2.4
-217.19	833.71	-7406.09	861.54	20465	4034	63.8	-3.6
-400.21	892.75	-7356.34	978.35	22966	4038	97.6	-2.5
-519.08	866.99	-7378.95	1010.5	20245	4042	58.8	-3.3
-112.1	1157.18	-7335.7	1162.6	22763	4047	45.7	-3.2
-764.24	497.39	-7425.14	911.84	23360	4051	103.9	-3
-737.72	222.25	-7334.98	770.47	19283	4066	87.4	-3.8
-1008.31	502.76	-7296.67	1126.7	19443	4066	42.7	-3.3
-333.11	1073.84	-7331.27	1124.32	23619	4098	92.8	-2.6
-338.44	1074.77	-7332.7	1126.8	23619	4098	-57.5	-2.6
-901.61	186.34	-7341.06	920.66	19543	4109	71.9	-2.9
-817.64	349.49	-7351.14	889.2	20109	4129	109.9	-3.6
-60.28	1362.16	-7328.29	1363.49	20346	4163	65.5	-1.4
-122.87	1258.53	-7351.81	1264.51	22316	4163	78.6	-1.2
-135.28	1337.62	-7348.75	1344.44	20535	4178	26.2	-2.9
-874.6	376.07	-7335.56	952.03	23502	4222	102	-2
-990.63	434.17	-7367.67	1081.6	20513	4222	31.2	-2.8
-65.3	936.13	-7369.22	938.4	20626	4264	67.5	-3.2
-940.41	95.92	-7342.78	945.29	20721	4294	70.5	-3.3
-271.21	1206.57	-7329.26	1236.68	19850	4311	27.7	-3.3
-731.8	538.3	-7329.25	908.46	20848	4324	71.3	-3.5
-905.21	236.82	-7349.22	935.68	19263	4327	67.6	-3.5
-694.34	-224.73	-7366.71	729.8	19340	4327	167.9	-4.2
-959.66	321.21	-7348.8	1011.99	19639	4337	60.9	-3.6
-794.49	551.59	-7294.59	967.19	23488	4351	104.7	-2.4
-713.44	567.22	-7323.26	911.45	23016	4351	67.7	-3.3
-720.47	548.62	-7323.73	905.57	23048	4371	109.2	-2.5
-901.32	240.41	-7327.97	932.83	21449	4371	68	-3.3

-291.91	922.1	-7489.48	967.2	22281	4383	61.2	-2.9
-459.72	896.67	-7333.29	1007.65	23589	4383	100.4	-2.8
-833.1	571.88	-7363.14	1010.5	22512	4392	101.1	-2.5
-423.53	959.15	-7346.84	1048.5	22769	4409	55.3	-3.2
-725.08	531.82	-7425.23	899.21	20880	4435	66.9	-3.5
-229.92	1110.04	-7404.21	1133.6	22035	4446	46.8	-3.1
-26.86	968.96	-7316.31	969.33	19285	4455	64.9	-3.2
-544.57	879.17	-7369.19	1034.16	22097	4482	58.2	-2.9
107.13	1415.43	-7383.54	1419.48	20570	4528	73.1	-2.8
-497.93	842.53	-7325.78	978.67	22231	4537	100.7	-2.7
-545.06	603.32	-7384.94	813.07	22099	4541	271.4	-3.2
-402.22	678.33	-7431.88	788.61	20482	4541	78	-3.6
-554.46	611.37	-7445.19	825.35	22127	4541	75.7	-3.2
-597.28	697.96	-7350.06	918.64	21591	4584	104.3	-3.1
-764.03	416.84	-7334.49	870.34	21402	4594	108.5	-3.2
-344.27	865.17	-7319.97	931.15	20610	4626	106.9	-3.1
-446.07	876.07	-7341.46	983.1	21968	4646	62.8	-3.1
-757.11	396.53	-7408.13	854.66	21279	4669	73.9	-3.4
-797.95	480.7	-7302.65	931.56	20588	4669	69.7	-3.6
99.21	938.23	-7320.33	943.46	21144	4675	69.4	-3.3
-207.17	1222.67	-7382.98	1240.1	22104	4679	38.1	-3.2
-392.56	811.04	-7468.79	901.05	21129	4688	68.8	-2.7
3.96	928.21	-7297.46	928.22	20800	4703	68.8	-3.3
-262.16	745.65	-7477.31	790.39	20474	4703	80	-3.5
-62.58	239.39	-7287.11	247.43	18724	4714	132.9	-5
-11.32	1190.87	-7301.9	1190.92	23076	4721	41.2	-3.1
-544.61	621.62	-7348.82	826.45	20893	4724	80.8	-3.4
-364.79	1040.21	-7306.94	1102.32	23088	4730	94.3	-2.7
489.9	1020.39	-7393.32	1131.9	21020	4734	43.6	-3.1
-561.83	763.28	-7357.61	947.76	23345	4737	106.7	-2.5
-222.9	1137.53	-7282.94	1159.16	20299	4737	40.9	-3
-104.12	1701.91	-7324.35	1705.09	22615	4753	53.1	-2.3
-1017.75	94.15	-7318.63	1022.1	23031	4771	98.7	-2.1
-432.14	843.52	-7340.77	947.77	23521	4779	102.8	-2.8
-225.17	1044.35	-7325.59	1068.35	22451	4789	58.7	-3.3
-568.54	932.32	-7331.11	1092	21614	4793	56.3	-3.1
-467.34	833.29	-7354.74	955.39	19291	4807	60.7	-3.6
289.05	1691.48	-7282.39	1716	21184	4814	55.2	-1.9
337.54	1677.48	-7288.09	1711.1	21232	4814	-95.8	-1.9
-808.12	524.08	-7383.52	963.18	19598	4862	59.9	-3.6
-296.39	387.06	-7391.96	487.51	19154	4871	107.9	-3.8
-401.77	535.98	-7390.2	669.85	19751	4871	98.4	-3.3
-276.92	-30.25	-7276.94	278.57	19130	4871	135.4	-4.4
-466.53	1142.38	-7362.05	1233.97	21475	4881	34.7	-2.9
-497.78	851.46	-7410.39	986.29	23527	4881	99.3	-2

-951.79	335.29	-7281.02	1009.12	22268	4887	99.7	-3.1
-938.72	-262.27	-7371.06	974.67	20952	4929	98.8	-2.8
-373.43	1016.81	-7334.72	1083.21	19101	4936	118.8	-2.8
-697.9	751.36	-7377.11	1025.48	18230	4936	289.5	-3.2
-666.77	630.6	-7342.08	917.74	21606	4965	107.4	-3
-343.35	1050.32	-7374.49	1105.02	23262	4977	92.4	-2.5
-270.21	1068.75	-7343.3	1102.38	20885	4986	53.1	-3.1
-259.17	350.28	-7290.51	435.74	20253	4994	347.4	-3.7
-435.41	943.71	-7367.54	1039.31	21981	4997	58	-2.9
-497.46	841.96	-7298.57	977.94	22692	5001	63.8	-3
361.68	1041.38	-7355.48	1102.4	19237	5021	49.3	-3.5
-351.21	946.08	-7391.73	1009.17	19042	5029	54.3	-3.2
-1052.78	82.82	-7302.98	1056.03	19626	5036	131	-2.6
-596.65	670.32	-7403.61	897.4	20984	5069	65.7	-3.5
-965.61	-29.26	-7393.78	966.05	21048	5073	64.1	-3.4
342.93	947.44	-7336.44	1007.59	19045	5082	62.7	-3.3
-365.12	744.34	-7416.07	829.07	22393	5082	322.8	-2.4
-719.8	585.86	-7383.69	928.09	21142	5115	98.8	-2.6
-719.16	514.08	-7392.03	884.01	20425	5124	68.4	-3.4
-903.79	259.1	-7339.43	940.2	22177	5135	104.1	-2.6
-670.05	89.47	-7282.41	676	21761	5139	122	-2.8
-272.48	1498.4	-7564.94	1522.97	22205	5139	-2.9	-2.9
-973.81	36.17	-7298.95	974.48	20073	5139	68.4	-3.3
-697.34	600.1	-7357.54	920	19481	5150	64.4	-3.6
-687.59	640.7	-7309.28	939.83	22374	5161	104.3	-2.8
233.91	1448.44	-7419.17	1467.21	21933	5171	11.1	-2.6
-863.61	327.43	-7343.85	923.6	23514	5178	106.3	-1.3
-409.18	1238.75	-7370.91	1304.58	22705	5222	80.3	-2.4
317.24	938.91	-7316.48	991.06	20979	5226	63.4	-3.1
-681.39	355.02	-7318.09	768.33	21095	5226	85.3	-3.1
-177.66	1145.12	-7425.47	1158.82	22578	5232	87.8	-2.2
-603.22	742.18	-7318.93	956.4	22412	5250	106	-2.5
-945.41	232.68	-7390.68	973.62	23393	5250	95.8	-3
-432.58	825.59	-7392.66	932.05	18065	5260	68	-3.5
-804.57	207.86	-7336.37	830.99	21136	5266	79.9	-3.4
-788.59	423.98	-7397.47	895.34	22685	5293	107.2	-3.2
-552.98	771.32	-7301.07	949.06	22661	5300	101.4	-3.7
-885.41	-383.64	-7352.29	964.95	22083	5304	59	-3.4
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-784.93	869.18	-7384.45	1171.15	20927	5312	89.5	-0.7
-627.26	655.82	-7333.29	907.5	21273	5318	108.6	-0.4
72.33	1204.22	-7355.46	1206.39	20276	5343	46.5	-3.2
-509.38	736.58	-7445.28	895.55	22244	5357	104	-3
-513.03	916.47	-7338.81	1050.29	21539	5394	53.9	-3.2

-290.86	701.27	-7538.21	759.2	22116	5397	79.7	-2.8
-499.31	779.18	-7318.05	925.44	20434	5403	101.1	-2.9
-282.07	1073.63	-7380.46	1110.07	21509	5403	53.9	-3.1
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-837.79	356.21	-7393.44	910.37	19482	5421	71.2	-3.3
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-717.63	739.84	-7370.96	1030.71	21964	5435	59.1	-3.1
-977.02	34.03	-7306.76	977.61	18517	5440	101.1	-3
-1034.92	371.25	-7369.87	1099.49	21336	5469	95.3	-2.9
44.09	1595.46	-7353.83	1596.07	20904	5495	-3	-3.2
-259.68	23.91	-7274.46	260.78	18342	5495	136	-5
-750.76	522.07	-7309.72	914.44	23472	5515	105.2	-2.8
-1002.8	243.52	-7415.9	1031.94	23384	5520	57.8	-2.9
-418.31	944.85	-7340.73	1033.31	22935	5525	96.9	-2.6
-1001.46	153.4	-7521.69	1013.14	21075	5533	50.2	-3.3
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-717.93	813.83	-7333.9	1085.24	22124	5556	85.1	-2.5
56.74	1608.02	-7319.25	1609.02	22733	5561	64.3	-2
63.97	1674.28	-7291.56	1675.5	21968	5561	-91.8	-2
-931.02	394.77	-7347.78	1011.26	20693	5566	61.7	-3.1
-277.56	1084	-7354.4	1118.97	22529	5570	49.4	-3
-357.56	628.65	-7344.61	723.22	19021	5576	88.6	-3.6
-535.73	562.35	-7375.29	776.69	18988	5596	80.5	-3.7
79.59	1193.94	-7378.72	1196.59	21830	5600	39.9	-3.2
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-902.08	216.01	-7363.93	927.58	22026	5627	69.2	-3.2
-705.34	803.5	-7427.62	1069.17	21129	5649	57	-3.3
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-339.25	1021.75	-7366.72	1076.6	19746	5693	54	-3.1
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208.4	1703.07	-7301.57	1715.77	22590	5737	48.4	-2.3
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-297.2	917.88	-7391.42	964.8	22494	5754	-8.3	-2.7
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-841.98	482.01	-7346.36	970.19	19895	5787	61	-3.4
-755.94	520.69	-7363.97	917.91	19795	5793	60.4	-3.9
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-399.97	1023.94	-7318.72	1099.29	22618	5808	94.5	-1.6
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127.27	1351.63	-7374.31	1357.61	22451	5901	19.4	-2.7
143.31	1616.05	-7440.76	1622.39	18530	5912	0.6	-2.4
-223.87	949.73	-7286.48	975.76	21477	5920	64.5	-3.5
-579.33	801.43	-7328.83	988.9	20391	5920	93.2	-3.1
30.25	1212.26	-7356.75	1212.64	19140	5937	29.2	-3.4
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-205.97	1103.6	-7354.79	1122.66	18621	5963	51.1	-3
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161.24	1672.4	-7343.5	1680.15	22599	5997	-6.9	-2.4
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-377.8	635.51	-7464.76	739.33	21027	6020	81.2	-3.5
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43.01	1251	-7359.91	1251.74	22127	6123	82.5	-2.8
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22.67	1712.95	-7296.41	1713.1	19760	6168	-94.4	-2.4
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169.72	1577.44	-7386.63	1586.54	21542	6203	-4.5	-2.7
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-606	563.7	-7324.17	827.64	23674	6222	113.7	-0.3
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-698.84	677.38	-7327.03	973.25	19762	6262	65.6	-3.6
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298.36	1288.49	-7323.02	1322.58	20636	6306	22.5	-3.4
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-553.16	805.89	-7326.95	977.47	23027	6357	-47.7	0.5
324.33	1528.29	-7262.1	1562.33	18848	6399	-62.4	-2.7
-611.45	740.94	-7366.85	960.66	23391	6415	103.3	-2.9
-851.05	190.76	-7333.57	872.17	18515	6432	76.3	-3.7
-849.58	182.64	-7283.85	868.99	18198	6451	75.6	-3.7
-1742.7	-828.25	-7373.9	1929.51	22752	6469	44.2	-2.5
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-150.54	1561.1	-7330.52	1568.34	22079	6478	65.5	-1.4
36.83	1572.62	-7334.6	1573.05	21899	6478	-85.3	-1.4
-243.02	918.95	-7290.49	950.54	19218	6501	312.8	-3.6
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-456.58	874.51	-7382.64	986.53	23633	6612	101.3	-2.5
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137.5	862.38	-7294.97	873.27	21794	6694	152.2	-3.3
28.61	1041.7	-7328.3	1042.09	21381	6705	57.8	-3.2
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381.52	1042.12	-7381.08	1109.76	21828	6719	48.1	-2.9
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107.7	1482.88	-7421	1486.79	22594	6793	75.6	-1.8
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-495.3	846.13	-7379.18	980.44	22192	7002	64.3	-3.4
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-665.35	332.36	-7317.83	743.74	21140	7041	65.5	-3.7
-919.12	30	-7313.34	919.61	22890	7041	102.4	-2.7
-1152.13	-101.92	-7316.56	1156.63	21876	7052	90	-2.3
-543.05	767.91	-7368.41	940.53	22490	7063	103.6	-2.8
-501.22	901.77	-7333.57	1031.7	23614	7066	95.9	-2.3
-565.27	736.19	-7474.61	928.17	22236	7066	98.7	-2.1
-573.55	837.57	-7377.08	1015.13	23715	7080	96.1	0.1
-515.38	777.61	-7387.4	932.9	23132	7105	107.2	-1.6
438.15	1023.47	-7369.38	1113.31	22652	7154	67.1	-2.6
-1150.58	-116.27	-7290.76	1156.44	22837	7163	91.1	-1.8
-1148.27	-116.12	-7291.36	1154.13	22838	7163	-58.9	-1.8
-508.62	198.4	-7337.96	545.95	19725	7173	105.7	-4.2
-372.91	791.65	-7440.98	875.08	23621	7178	72.8	-2.7
-431.48	636.93	-7477.28	769.32	23441	7188	114.4	-1.2
-997.58	40.46	-7306.73	998.4	22992	7204	103.8	-1.7
-249.95	1017.73	-7381.93	1047.97	22106	7212	100.6	-2.6
-416.49	956.79	-7379.51	1043.51	22031	7218	56	-3.3
-402.22	1033.28	-7329.73	1108.81	22628	7218	52.9	-2.7
-524.8	878.92	-7426.33	1023.68	20653	7227	60.4	-3
-293.53	1147.81	-7322.83	1184.75	23415	7235	43.7	-2.8
-611.28	793.2	-7342.83	1001.41	23067	7235	100.6	-2.7
-540.69	839.31	-7306.75	998.39	23656	7249	100.5	-2.3
-994.78	-156.83	-7393.56	1007.07	21625	7249	61.4	-2.3
-504.38	879.44	-7355	1013.81	22750	7262	61.3	-3.1
-441.2	1009.29	-7414.79	1101.51	23023	7262	46.6	-2.8
-536.78	776.89	-7328.58	944.29	23624	7268	106.1	0.5
-727.26	524.82	-7418.58	896.85	22024	7279	106.7	-2
-680.6	623.95	-7331.09	923.33	20778	7279	-9.1	-3.1
-497.01	797.8	-7321.98	939.95	23002	7297	106.8	-2.5
-553.01	812.17	-7336.65	982.57	23612	7311	103.3	-0.4
-662.09	776.92	-7381.55	1020.77	22955	7316	57.8	-1.6
-675.35	636.77	-7403.91	928.21	23048	7316	105.2	-1.3
-899.73	95.78	-7318.99	904.81	23562	7316	-20.3	-2.6
-186.98	127.56	-7287.03	226.35	20506	7347	253.8	-4.5
146.39	518.32	-7366.54	538.6	20089	7353	129.8	-3.9
-752	619.49	-7426.67	974.31	23645	7359	272.9	-2.4
-499.41	834.14	-7422.3	972.21	22894	7359	101.1	-2.4
-913.47	289.36	-7329.1	958.2	22057	7363	104	-1.2
-347.18	655.4	-7491.03	741.68	20790	7382	81.7	-3.1

-538.47	774.94	-7329.94	943.65	23230	7389	104.6	-0.7
-605.37	749.91	-7459.7	963.76	22203	7410	210.3	-2.9
-604.23	759.38	-7455.88	970.44	22216	7410	61.8	-2.9
-301.54	937.62	-7396.39	984.92	22999	7418	63.1	-2.9
-803.39	470.68	-7295.73	931.11	19275	7423	67.2	-3.4
-584.82	690.42	-7331.11	904.82	19918	7438	68.9	-3.3
-549.27	883.11	-7331.13	1039.99	20249	7455	60.2	-3
162.16	1586.05	-7343.04	1594.32	22019	7459	-2.6	-2.9
63.93	1504.57	-7357.8	1505.93	19891	7471	4.6	-3.2
-963.32	165.49	-7390.41	977.43	19464	7476	64.1	-3
-479.49	880.43	-7344.27	1002.53	19388	7500	62.3	-3.1
-554.67	809.02	-7334.08	980.9	23635	7508	103.8	-0.8
-756.71	-365.44	-7327.72	840.33	21425	7535	78.1	-3.2
-450.73	910.58	-7359.01	1016.03	20105	7546	61.4	-3.5
-12.12	221.4	-7276.33	221.73	19918	7546	91.9	-5.1
-898.87	306.92	-7425.32	949.82	19616	7558	100.2	-3.5
120.54	1314.27	-7344.82	1319.79	22397	7570	79.9	-2.1
-561.84	918.31	-7357.03	1076.55	21696	7599	56.2	-3
-667.1	650.6	-7374.02	931.83	23714	7611	107.1	-1.8
-810.95	471.02	-7379.04	937.82	21637	7611	161.3	-1.7
-666.58	651.21	-7374.47	931.88	23699	7611	-42.9	-1.8
-811.27	467.87	-7381.96	936.52	21640	7611	11.3	-1.7
-25.7	1556.06	-7384.18	1556.27	22339	7614	64.7	-2.4
-391.8	1008.11	-7304.5	1081.57	23711	7622	95.3	-2.4
-631.12	732.91	-7318.94	967.2	23178	7633	97.4	-2.6
-585.51	747.54	-7335.26	949.55	22897	7637	106.2	-2.9
-1004.68	19.5	-7408.56	1004.87	22882	7644	54.8	-3.2
-377.3	1171.22	-7329.14	1230.49	19379	7644	39.2	-3.1
-342.19	927.77	-7422.34	988.86	22184	7649	93.7	-3.4
-75.83	1664.96	-7413.24	1666.69	21312	7667	-7.8	-2.8
-660.21	613.9	-7327.09	901.53	19823	7679	74.3	-4
-267.47	-7.58	-7284.17	267.58	19956	7684	138	-4.9
-1004.58	252.26	-7322.09	1035.77	19839	7691	57.9	-3.6
-1149.74	-255.34	-7347.17	1177.75	20661	7710	89.5	-3.2
-1079.88	-49.57	-7354.88	1081.02	23087	7725	97	-0.1
-273.74	1106.23	-7322.25	1139.6	21824	7760	45.8	-3.3
-77.65	237.22	-7282.36	249.61	18589	7760	405.3	-4.9
-421.05	877.32	-7364.08	973.13	20816	7772	61.4	-3.2
-20.37	452.78	-7350.56	453.24	20285	7772	119.7	-3.9
233.69	1720.15	-7293.88	1735.95	21534	7775	56.5	-1.6
-453.28	823.17	-7378.22	939.72	23600	7779	102	-2.5
-901.66	40.81	-7469.06	902.58	21673	7804	66.6	-3.4
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-624.02	654.49	-7329.57	904.3	23714	7859	72.5	-3
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-575.07	744.2	-7351.36	940.5	18867	7877	367.7	-3.8
-273.72	50.71	-7285.13	278.38	19740	7892	133.1	-4.6
-907.26	402.71	-7388.4	992.62	23093	7914	98.9	-2.9
-935.85	106.59	-7440.48	941.9	22659	7930	328.1	-3.5
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-1367.73	-216.27	-7299.46	1384.72	20593	7998	76	-3.3
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98.6	911.87	-7379.74	917.19	21728	8054	104.7	-3.1
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326.6	1322.89	-7337.86	1362.61	19744	8105	25.1	-2.9
301.65	1826.46	-7258.02	1851.2	22216	8116	49.5	-1
228.9	1906.65	-7333.05	1920.34	22234	8116	-109.1	-1
-361.23	828.87	-7323.06	904.16	19265	8129	71.3	-3.4
-589.96	763.62	-7357.93	964.97	21612	8129	66.6	-2.9
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-847.53	366.38	-7337.35	923.33	20668	8170	321.5	-3.5
-937.52	305.42	-7320.22	986.01	18832	8170	64.5	-3.4
-848.63	363.63	-7338.94	923.25	21152	8170	105	-3.5
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44.66	1101.5	-7282.42	1102.41	20668	8177	48.5	-3
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-600.47	670.9	-7358.78	900.37	21569	8208	104.1	-3.3
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-924.45	95.13	-7326.31	929.33	23709	8220	107.7	-0.5
-898.11	371.73	-7346.6	972	20730	8243	134.5	-2.7
-1135.64	-82.83	-7268.6	1138.66	23279	8252	94.4	-1.1
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-483.19	859.87	-7316.03	986.33	21539	8277	-53.5	-2.6
344.61	977.08	-7332.45	1036.07	19525	8286	60.4	-3.4
-810.51	296.95	-7318.92	863.2	23544	8317	111	-0.9
-908.89	-276.71	-7399.58	950.08	20262	8327	65	-3.5
-404.08	903.34	-7357.2	989.6	22986	8330	101	-2.7
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-572.08	765.65	-7330.41	955.77	23480	8337	105.4	-2.1
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-550.79	782.37	-7331.11	956.8	22927	8365	102	-2.2
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57.94	1245.36	-7378.8	1246.71	21906	8708	30.8	-3.1
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-33	1177.01	-7364.3	1177.47	21125	8792	44.1	-3
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-103.11	1212.42	-7452.94	1216.8	19329	9015	35.6	-3.5
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-617.7	717.38	-7398.07	946.67	23629	9094	99.5	-3
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-84.27	1224.54	-7361.05	1227.44	19363	9255	31.5	-3.4
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45.75	1128.53	-7373.36	1129.46	18289	9432	230.3	-3.5
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-76.16	1177.34	-7372.64	1179.8	22993	9514	43.2	-2.5
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-392.54	-58.7	-7279.03	396.9	20560	9575	125	-4.5
-1070.14	157.04	-7355.47	1081.6	22860	9607	95.1	-2.3
-231.1	1243.05	-7337.19	1264.35	22555	9676	82.8	-2.3
-372.13	572.7	-7339.12	682.98	22055	9690	93.7	-3.8
-991.04	14.74	-7322.37	991.15	22478	9695	66.5	-2.2
-836.08	92.36	-7286.57	841.17	21022	9700	112.2	-3.1
-1148.03	-120.96	-7331.14	1154.38	22055	9712	90	-2
-1274.74	-146.45	-7327.23	1283.12	23010	9722	80	-1.5
-1142.43	-73.39	-7378.41	1144.78	23471	9741	50.5	-2.8
-1093.14	-139.64	-7316.88	1102.02	21993	9769	95.2	-1.6
-1051.96	98.44	-7380.81	1056.56	20555	9788	47.9	-3.2
-336.24	1150.53	-7368.46	1198.66	23478	9801	87.8	-2.5
-689.38	636.11	-7317.67	938.02	23170	9806	105.3	-1.6
-664.62	595.8	-7325.16	892.58	23805	9810	109.3	-0.5
-926.21	-138.2	-7317.08	936.46	21851	9845	106.8	-1.8
-708.94	651.62	-7321.38	962.91	23210	9850	105.8	-2.4
75.8	1464.91	-7394.42	1466.87	22299	9853	14.1	-2.8
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-570.08	623.45	-7354.26	844.8	21592	9892	66	-3.6
-872.2	300.59	-7338.96	922.54	19361	9916	68	-3.6
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-613.47	822.01	-7360.23	1025.69	22851	9946	55.5	-2.8
-1005.47	-47.18	-7336.55	1006.58	21206	9954	93	-3.1
-718.25	819.71	-7374.96	1089.87	21152	9960	51	-3.3
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-490.03	857.4	-7363.11	987.55	23056	9996	63.8	-3.3
-267.98	1108.76	-7327.55	1140.68	21926	10005	47.4	-3.3
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-211.19	-197.17	-7413.18	288.92	21800	10050	139.2	-4.8
-995.99	-122.86	-7315.29	1003.54	20374	10050	62.4	-3.4
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329.11	1139	-7355.46	1185.59	20598	10080	38.3	-3.2

-432.74	974.02	-7329.67	1065.82	20885	10088	54.2	-3
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-65.16	1212.85	-7323.79	1214.6	19932	10183	39	-3.5
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14.77	1412.09	-7325.89	1412.17	19274	10211	73.7	-1.8
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142.88	1109.74	-7408.2	1118.9	22865	10239	83.4	-2.9
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-845.94	-153.44	-7332.34	859.74	18642	10255	71.8	-3.6
-816.9	482.82	-7412.36	948.92	22477	10287	64.3	-3.4
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-799.26	550.7	-7382.26	970.61	21466	10496	100.5	-3.5
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-888.74	-95.67	-7414.41	893.87	20937	10556	67.4	-3.5
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-1015.9	-701.24	-7358.79	1234.42	22512	10636	76.6	-2.5
77.47	1528.79	-7362.51	1530.75	22344	10639	59.2	-2.8
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163.94	1923.15	-7387.18	1930.12	21089	10712	40.8	-2.4
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410.9	1666.07	-7331.13	1715.99	23135	10821	56.6	-1.9
374.29	1578.64	-7331.11	1622.4	23136	10821	-88.7	-2
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-353.8	791.72	-7325.65	867.18	19497	10966	73.7	-3.5
-468.62	894.92	-7360.01	1010.19	23076	10974	92.1	-2.9
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102.55	1087.33	-7368.22	1092.16	18443	11264	53	-3
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40.73	761.62	-7297.45	762.71	20035	11390	86.6	-3.2
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5.97	1401.34	-7383.04	1401.35	20833	12148	18.4	-2.8
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22.36	1170.9	-7420.33	1171.11	20527	12211	38.9	-3.4
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-980.46	51.29	-7295.98	981.8	22051	13753	103.6	-2.9
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-993.73	-90.52	-7430.95	997.84	23584	13814	99.5	-2.6
-761.44	449.11	-7306.65	884.02	22707	13814	71.5	-2.8
-893.02	63.6	-7295.49	895.28	23331	13829	108.6	-1.9
-230.69	563.82	-7292.6	609.19	19763	13836	102.3	-3.7
-875.89	82.88	-7300.47	879.8	21697	13842	76.7	-3
-846.07	57.63	-7352.04	848.03	18650	13855	77.6	-3.9
-233.08	1205.18	-7368.7	1227.51	19960	13930	36.2	-3.3
-270.09	983.33	-7295.4	1019.75	21841	13966	59.8	-2.7
-301.65	1307.01	-7413.16	1341.37	20415	13966	27.9	-3.1
-574.06	912.37	-7394.9	1077.94	21113	13982	45.2	-3.6
-207.56	1549.77	-7338.58	1563.61	20074	14037	58.3	-2.7
-33.72	1497.18	-7343.37	1497.56	22919	14054	64.5	-2
-999.59	143.62	-7294.64	1009.85	23576	14062	62.5	-2.6
26.45	2291.96	-7231.92	2292.11	23345	14066	-62.1	-2.4
-73.21	1472.65	-7339.01	1474.47	22634	14076	69	-0.9
-68.16	1485.63	-7343.29	1487.19	23017	14076	-80.7	-0.9
-120.2	1459.36	-7334.64	1464.3	21645	14076	8.7	-1.9
-131.19	1403.46	-7339.49	1409.58	21835	14090	75.5	-2.1
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-141.57	1397	-7347.3	1404.15	21353	14237	74.1	-0.9
-119.81	1131.63	-7473.45	1137.95	21300	14237	-55.9	-1.1
-95.81	1248.81	-7396.32	1252.48	20924	14246	35.3	-2.5
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117.51	1367.48	-7423.12	1372.52	21925	14451	20.9	-2.9
-79.03	1396.2	-7381.21	1398.43	22289	14451	23.1	-2.5
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-232.95	1079.2	-7379.29	1104.06	22496	14793	48.4	-2.9
-536.66	365.56	-7289.5	649.34	23193	14793	98.5	-3.7
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-451.12	462.89	-7297.2	646.36	23186	14824	125.1	-2.3
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-307.29	1190.04	-7352.56	1229.07	19845	14949	39.7	-3
64.54	251.86	-7355.48	260	20454	14968	138.7	-4.7
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-208.33	963.19	-7341	985.46	22126	15902	62.5	-3.5
-902.46	-165.37	-7342.47	917.49	23387	15909	109	-1.5
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-52.92	1204.53	-7362.19	1205.69	21586	15943	40.3	-3
339.81	1180.78	-7391.84	1228.7	19645	15947	33.4	-3.6
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-796.17	-215.44	-7381.44	824.8	21701	16035	108.4	-3
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-951.41	-149.04	-7363.02	963.01	20662	16051	65.1	-3.6
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-457.38	603.4	-7330.7	757.16	21025	16124	108.8	-3.1
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192.23	1600.5	-7440.76	1612	23099	18422	-2.1	-1.9
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291.68	1955.08	-7211.7	1976.72	23625	21122	40.6	-1.9
69.97	1932.09	-7318.83	1933.36	23662	21122	-109.2	-1.9
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-411.37	753.92	-7296.21	858.85	23624	21238	107.9	-3
-260.12	877.45	-7294.56	915.19	20427	21284	70.7	-3.6
-986.25	-279.82	-7356.22	1025.18	22927	21311	97.9	-2.9
-1373.03	-225.14	-7353.08	1391.37	18673	21316	23.2	-3.2
-72.48	1442.35	-7346.97	1444.17	21883	21337	14.2	-2.8
-887.18	-249.84	-7355.85	921.69	21403	21457	64.8	-3.9
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11.91	1565.87	-7344.2	1565.92	22311	21553	64.7	-1.1
-1004.18	-131.21	-7376.09	1012.72	21646	21623	60.4	-3.7
-492.69	1114.51	-7392.24	1218.55	19506	21665	-43.2	-3.5
-452.98	929.34	-7135.93	1033.86	23358	21686	64.6	-3
-473.13	-369.2	-7314.06	600.13	19308	21698	124.3	-3.6
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-663.64	-209.67	-7394.17	695.97	21671	21722	120.8	-3.1
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-1518.33	178.55	-7355.46	1528.79	23116	21828	63	-2.2
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6.03	1183.5	-7360.35	1183.52	23554	22255	85	-2.6
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-196.92	1023.93	-7130.11	1042.69	23423	22524	99.6	-2.8

-171.56	1018.13	-7410.16	1032.48	21089	22608	53.6	-3.2
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-298.74	650.26	-7353.1	715.6	23602	22673	119.7	-2.9
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79	744.6	-7306.77	748.78	20987	22725	87.4	-4.4
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-754.87	-220.7	-7388.76	786.47	23475	23600	116.4	-0.8
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-434.82	757.69	-7306.79	873.59	23287	23943	107.6	-3.3
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161.96	821.83	-7346.35	837.64	19152	24158	40.8	-4
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283.85	2384.98	-7188.98	2401.81	22276	24534	15.7	-0.7
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237.79	559.1	-7287.65	607.57	23695	25793	125.7	-3
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32.45	1695.8	-7358.04	1696.11	23116	25919	-98.9	-1.4
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-769.4	-233.18	-7375.49	803.96	23090	31855	114.7	-0.4
72.86	462.08	-7370.7	467.79	21234	31926	113.7	-4.7
-872.67	-105.33	-7470.38	879	23423	31931	70.7	-3.4
-1204.96	-383.28	-7348.02	1264.45	23155	32004	77.2	-2.7
-1116.73	-244.65	-7340.55	1143.21	23178	32004	-64.3	-2.8
138.29	1284.23	-7321.05	1291.65	22926	32008	35.4	-3.3
376.08	1329.9	-7405	1382.05	22143	32011	21.3	-3.4
176.72	1489.74	-7317.39	1500.19	21067	32011	66.9	-3.2
-881.36	-264.95	-7348.1	920.32	21732	32053	109.4	-3
-765.08	-239.68	-7433.8	801.74	23134	32068	113.6	-0.8
-362.09	671.95	-7438.9	763.3	22128	32140	84.7	-4.1
-413.07	671.02	-7300.16	787.97	18849	32184	78	-4.5
71.83	1366.42	-7361.92	1368.31	20333	32238	24.1	-3.4
116.26	820.1	-7379.04	828.3	19741	32254	52	-4.4
-492.91	-78.98	-7294.59	499.2	19646	32310	116.7	-4.3
-963.3	-53.21	-7371.85	964.77	21599	32361	61.7	-4.1
-21.52	686.07	-7331.13	686.41	21471	32409	94.6	-3.9
99.15	928.78	-7524.36	934.06	20278	32519	55.1	-4.2
-73.52	401.52	-7364.13	408.2	22003	32526	139.6	-3.8
16.69	343.59	-7319.36	344	19159	32558	131.2	-4.6
-777.53	0.51	-7349.86	777.53	22613	32580	115.7	-3.4
-523.67	-277.82	-7282.36	592.8	20369	32686	128.4	-4.1
-809.15	-157.68	-7386.18	824.37	20653	32787	112.8	-4.3
-194.85	390.95	-7331.11	436.82	20439	32887	140.9	-2.5
-268.71	459.41	-7344.47	532.22	21845	32890	133.3	-3.8
-114.37	384.68	-7376.28	401.32	22739	33039	143.5	-4.7
-281.24	1011.94	-7315.17	1050.29	23445	33112	98.6	-1.2
-279.04	1005.04	-7314	1043.06	23454	33112	-51.7	-1.2
77.93	1256.04	-7371.24	1258.46	23297	33145	82.7	-2.7
153.11	1342.53	-7348.22	1351.23	23615	33154	73.5	-3.2
-705.6	165.46	-7349.9	724.74	20165	33258	84.6	-4.5
207.47	1420.71	-7428.24	1435.78	21876	33292	15.9	-3.4
-46.05	1306.6	-7365.82	1307.41	18992	33292	33.1	-3

-295.94	-127.91	-7294.58	322.4	20523	33330	140	-4.2
-181.92	487.73	-7290.38	520.55	22392	33339	395.3	-4.6
-171.49	468.73	-7288.45	499.12	21866	33339	115.1	-4.7
-538.8	568.24	-7241.31	783.07	18172	33403	85.8	-4.3
168.59	1305.62	-7374.04	1316.46	22533	33409	28.6	-3.4
148.44	1349.95	-7336.69	1358.09	22269	33677	74.1	-3.5
-150.01	385.76	-7388.78	413.9	19551	33707	109.1	-5.2
31.18	1388.99	-7334.59	1389.34	18542	33732	17.2	-3.7
-332	167.89	-7335.19	372.04	20643	33767	138.6	-3.7
-332.84	161.82	-7350.02	370.09	21760	33778	148.2	-2.9
-323.89	130.31	-7329.96	349.12	19368	33806	145.1	-4.1
-267.73	543.73	-7329.08	606.07	22577	33822	126.8	-3
-307.32	684.56	-7374.18	750.38	21614	33905	118	-3.4
76.58	1310.2	-7389.31	1312.44	19372	33919	23	-3.7
-287.78	63.15	-7229.18	294.63	19350	33942	134.4	-5.3
34.51	974.19	-7372.97	974.8	20689	34201	100.2	-3.6
69.5	963.91	-7328.58	966.41	18101	34231	64.9	-4.2
-702.29	108.98	-7329.15	710.7	18883	34313	92.5	-4.2
-173.28	149.45	-7319	228.83	20622	34342	155.1	-5
-135.42	188.81	-7342.5	232.35	22957	34472	349.8	-4.8
-157.53	213.44	-7350	265.28	22153	34477	151.1	-4
-121.21	185.7	-7333.31	221.76	21898	34477	148	-5
-364.79	-288.27	-7288.51	464.94	18325	34511	117.5	-5.1
-362.87	682.56	-7354.34	773.02	23109	34630	116.3	-3.5
-440.78	592.24	-7340.38	738.27	20942	34634	90.4	-4.3
-832.62	-224.36	-7461.94	862.32	23332	34694	103	-2.9
242.18	1297.71	-7375.63	1320.11	20891	34727	29.6	-3.5
-465.39	-182.36	-7527.65	499.84	21321	34786	130.7	-4.4
-1187.94	-349.46	-7323.55	1238.27	22823	34843	83.1	-2.5
-205.68	84.58	-7285.57	222.39	18119	35033	238.6	-5.5
418.51	1425.82	-7496.47	1485.97	19270	35119	-6	-3.7
-194.69	-138.97	-7343.32	239.2	19941	35148	143.9	-5.2
-215.19	1031.71	-7227.33	1053.91	23503	35156	60.9	-3
-312.91	597.27	-7326.52	674.27	21063	35178	95.3	-4.2
-696.98	-140.1	-7473.88	710.92	21116	35195	109.7	-3.4
-301.37	546.14	-7289.05	623.77	22437	35203	121.5	-4
-862.9	-243.88	-7314.42	896.7	22536	35313	108.4	-3.5
93.03	1356.88	-7379.97	1360.07	19088	35382	-16.5	-4
-173.38	1009.29	-7259.69	1024.07	22867	35510	60.9	-3.6
-187.15	829.64	-7347.05	850.49	23589	35513	78.4	-3.3
-35.35	969.45	-7422.33	970.09	22666	35630	101.8	-3.8
94.76	1279.55	-7364.14	1283.05	22085	35876	33.2	-3.6
9.43	316.8	-7275.26	316.94	18058	35902	147.4	-4
450.54	1194.23	-7356.96	1276.39	19487	35911	35.4	-3.7
-211.93	629.65	-7285.36	664.36	21061	35998	95.4	-4.2

-127.83	610.78	-7297.97	624.01	21137	36159	102.9	-3.8
-186.45	-90.76	-7403.63	207.37	21535	36202	149.9	-4
155.08	1233.52	-7389.49	1243.23	19213	36246	35	-3.8
-46.03	903.63	-7355.47	904.8	22093	36347	74.5	-3.7
-135.54	651.2	-7317.64	665.16	23600	36570	123.6	-3.4
-363.29	671.22	-7368.69	763.23	22275	36667	113.9	-3.9
-101.65	613.62	-7287.9	621.98	21295	36694	101.1	-4.5
-947.1	495.72	-7388.57	1068.99	18250	36711	52.9	-4
6.22	351.66	-7328.61	351.72	19156	36753	132.1	-1.6
-239.65	-108.98	-7361.59	263.27	23066	36898	140.7	-4.2
-202.33	224.46	-7338.36	302.19	22634	36903	148.7	-4.3
-127.18	652.6	-7302	664.88	21078	36927	93.4	-4.1
-553.09	-228.26	-7313.74	598.34	21424	36951	127	-3.7
-498.2	-194.52	-7359.56	534.83	20411	37087	128.1	-3
-848.13	-246.25	-7357.58	883.16	22403	37090	110	-1.9
-432.2	-242.01	-7338.46	495.34	21872	37209	118.3	-4.3
-773.36	-214.85	-7375.78	802.65	23742	37236	114.7	-2.1
-775.2	-246.53	-7359.34	813.46	23651	37353	117	-2
-759.56	-309.45	-7348.07	820.18	20485	37379	113.5	-4.3
-32.8	920.48	-7256.31	921.06	21333	37402	71.8	-2.5
-743.97	803.12	-7317.98	1094.76	20437	37632	95.3	-3.7
-84.5	357.97	-7326.81	367.81	19759	37820	130.9	-1.3
-247.03	733.98	-7346.65	774.44	21322	37852	89.5	-4.1
-92.48	620.45	-7323.49	627.3	23028	37875	107.3	-4
-23.56	1258.17	-7428.56	1258.39	20123	37912	29.5	-3.7
37.31	799.82	-7351.92	800.69	19350	37925	85.1	-4
63.16	750.73	-7332.13	753.38	20904	37925	117.8	-4.3
-12.68	776.72	-7335.58	776.82	19280	37930	87.5	-4.3
63.37	830.33	-7353.76	832.74	20572	37990	80.7	-4.2
-127.25	680.77	-7331.06	692.56	20226	38012	121	-4.4
-749.62	-50.93	-7301.52	751.35	21916	38321	88.9	-3.9
-1229.52	-148.22	-7493.05	1238.42	22081	38399	82.8	-2.7
-1022.62	-57	-7213.79	1024.21	23421	38732	99.3	-3.2
-58.5	608.88	-7347.01	611.68	22727	17274	128.1	-4.3
-314.38	-144.25	-7377.33	345.89	19435	17601	131.3	-4.1
-544.93	761	-7306.79	935.99	21577	17616	65	-3.9
-267.01	529.24	-7318.96	592.78	18839	17637	121.3	-4.7
-491.76	-173.56	-7359.02	521.49	21489	17794	131.7	-3.9
-710.6	-7.78	-7318.52	710.64	21484	17858	117.8	-2.2
-156.29	385.53	-7270.18	416	21711	17913	125.2	-3.2
-230.39	494.41	-7290.5	545.45	21676	17923	109.2	-3.9
-199.43	-62.58	-7341.56	209.02	20401	17923	158.1	-4.4
440.25	1409.64	-7465.11	1476.79	21119	18134	74.1	-3
-240.01	428.01	-7280.22	490.71	23186	18152	137.2	-4.3
-207	-85.88	-7329.96	224.11	21559	18227	156.1	-2.5

-311.04	-45.15	-7280.54	314.3	18930	18669	138	-3.6
-507.75	38.95	-7304.07	509.24	23568	18964	135.1	-2.4
-200.14	777.71	-7317.99	803.05	21148	19048	84.5	-4.2
-261.22	235.31	-7330.77	351.58	20421	19063	132.4	-5
-157.79	340.94	-7393.21	375.68	21093	19152	142.6	-4.3
-542.13	556.07	-7285.7	776.61	21071	19172	81.7	-4
-358.9	331.84	-7282.41	488.8	22259	19343	114.5	-4.7
-45.05	316.23	-7288.38	319.42	19139	41278	134.4	-5.1
98.68	1688.32	-7222.21	1691.2	22375	41597	59	-1.9
-294.73	-189.02	-7369.49	350.13	18066	41616	119.3	-5.4
-10.38	520.05	-7323.23	520.15	23138	41828	136.1	-2.9
362.61	1403.56	-7351.33	1449.64	21587	41974	13.7	-3.3
-549.9	909.55	-7381.54	1062.86	18636	42168	56.6	-4
-432.91	-144.96	-7283.58	456.54	21317	42551	135.8	-3.7
-594.32	473.79	-7272.6	760.06	23121	42573	116.4	-3.4
113.07	527.3	-7365.09	539.29	21004	42608	133	-4.6
632.22	1527.98	-7270.22	1653.61	22120	42879	57.8	-2.7
337.75	1424.88	-7400.18	1464.36	20537	42885	14.8	-3.5
364.76	1555.59	-7249.77	1597.78	22267	43028	1.7	-3
-331.85	189.08	-7410.52	381.94	21448	43303	144.8	-4.7
-49.37	318.6	-7282.37	322.4	21810	43740	148.1	-2.9
-470.31	-173.44	-7279.29	501.27	21103	43830	132.6	-4.1
147.47	478.38	-7348.64	500.59	20885	44270	117.4	-4.3
-737.47	559.36	-7332.32	925.61	23121	44281	108	-3.3
-774.25	510.44	-7327.3	927.37	20854	44356	107.6	-3.7
-62.71	326.69	-7329.37	332.65	18205	44799	145.7	-4.3
182.22	533.31	-7329.37	563.58	22338	44910	131.7	-3.6
-341.79	-299.36	-7342.53	454.35	19861	45050	121.8	-4.5
-517.21	56.8	-7296.81	520.32	18064	45107	109.8	-4.8
-435.8	-223.02	-7288.46	489.55	18993	45253	112.5	-4.8
-570.05	14.89	-7326.44	570.24	19118	45562	109.5	-4
-773.26	-133.86	-7299.22	784.76	22197	45709	110.5	-3.3
-96.96	1249.3	-7355.56	1253.06	18854	45984	35	-3.6
-458.53	266.75	-7293.38	530.48	23019	46343	115.1	-4.2
-289.28	-184.67	-7343.31	343.2	22476	46701	137.7	-3.2
379.14	1171.17	-7385.06	1231.01	18866	46930	33.6	-3.5
-322.66	-146.44	-7302.21	354.34	19397	47415	132.7	-4.4
-232.93	-189.53	-7342.9	300.3	19658	48084	134.8	-5.3
-320.66	-89.1	-7295.24	332.81	19038	48592	127.3	-4.9
-335.12	-111.61	-7341.98	353.22	19624	49852	131.7	-4.3
6.55	1550.54	-7344.54	1550.55	23658	50489	67	-2
-119.84	1469.04	-7377.81	1473.92	20812	51161	70.5	-2.2
-119.88	1466.49	-7376.1	1471.38	20817	51161	138.2	-2.2
-648.89	503.28	-7372.92	821.19	19916	53803	78.6	-3.4
-318.89	41.36	-7370.49	321.56	19703	54458	133.7	-4.9

-543.76	852.67	-7281.94	1011.3	19232	55986	56.6	-3.8
-327.74	300.7	-7293.79	444.79	21512	56886	137.8	-3.6
-796.91	501	-7337.97	941.31	18826	57111	71.8	-4
-98.51	407.89	-7311.02	419.62	22895	57345	144	-3.8
-119.03	396.32	-7291.87	413.81	21489	57550	143.5	-3.4
-241.9	-52.58	-7338.93	247.55	22721	62018	143.2	-3.8
-523.32	953.2	-7345.49	1087.41	20688	63064	94.3	-3.7
-280.22	-152.72	-7295.66	319.13	21151	63409	146.6	-3.7
-232.71	-18.03	-7373.42	233.41	20219	63667	145	-5
-34.46	406.01	-7295.64	407.47	23028	64543	142.2	-3.2
-4.17	393.61	-7282.91	393.63	23321	64547	145.2	-2.8
1.53	365.23	-7295.81	365.23	20985	64758	145.4	-4.4
45.29	286.17	-7328.09	289.73	20126	66710	139.7	-2.2
-565	781.52	-7323.14	964.36	22616	68485	104.3	-2.3
-229.59	936.65	-7323.03	964.38	22606	69264	101.7	-3.7

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